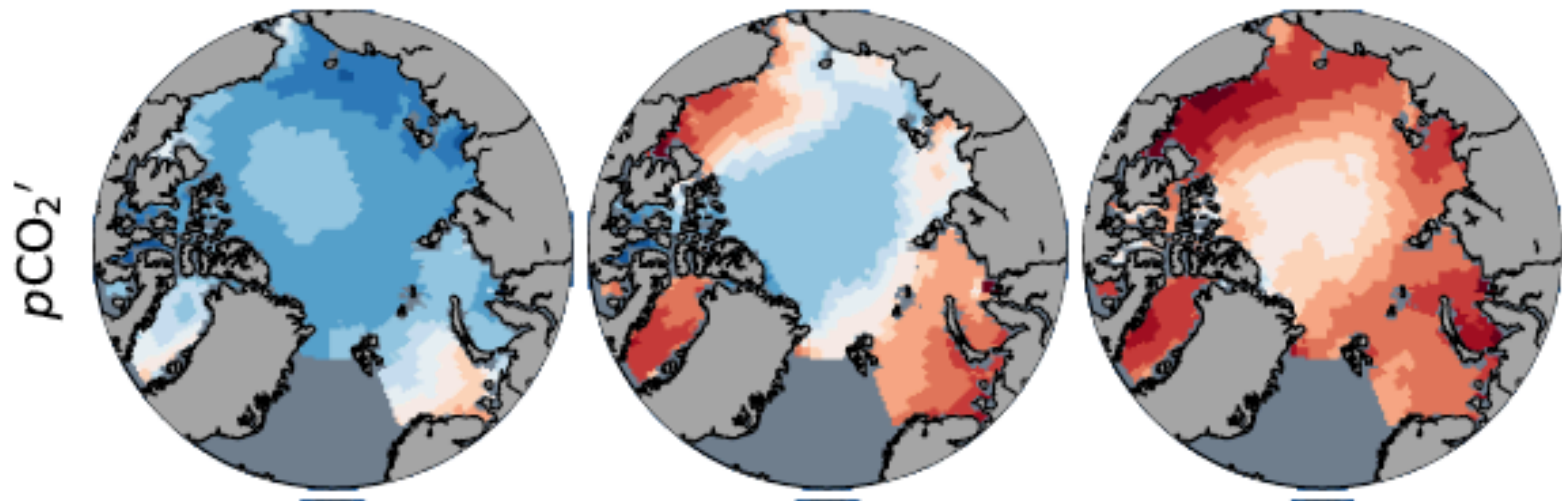


# Ocean biogeochemical tipping points in ESMs: a regime shift in Arctic ocean chemistry ?



Lester Kwiatkowski & James Orr

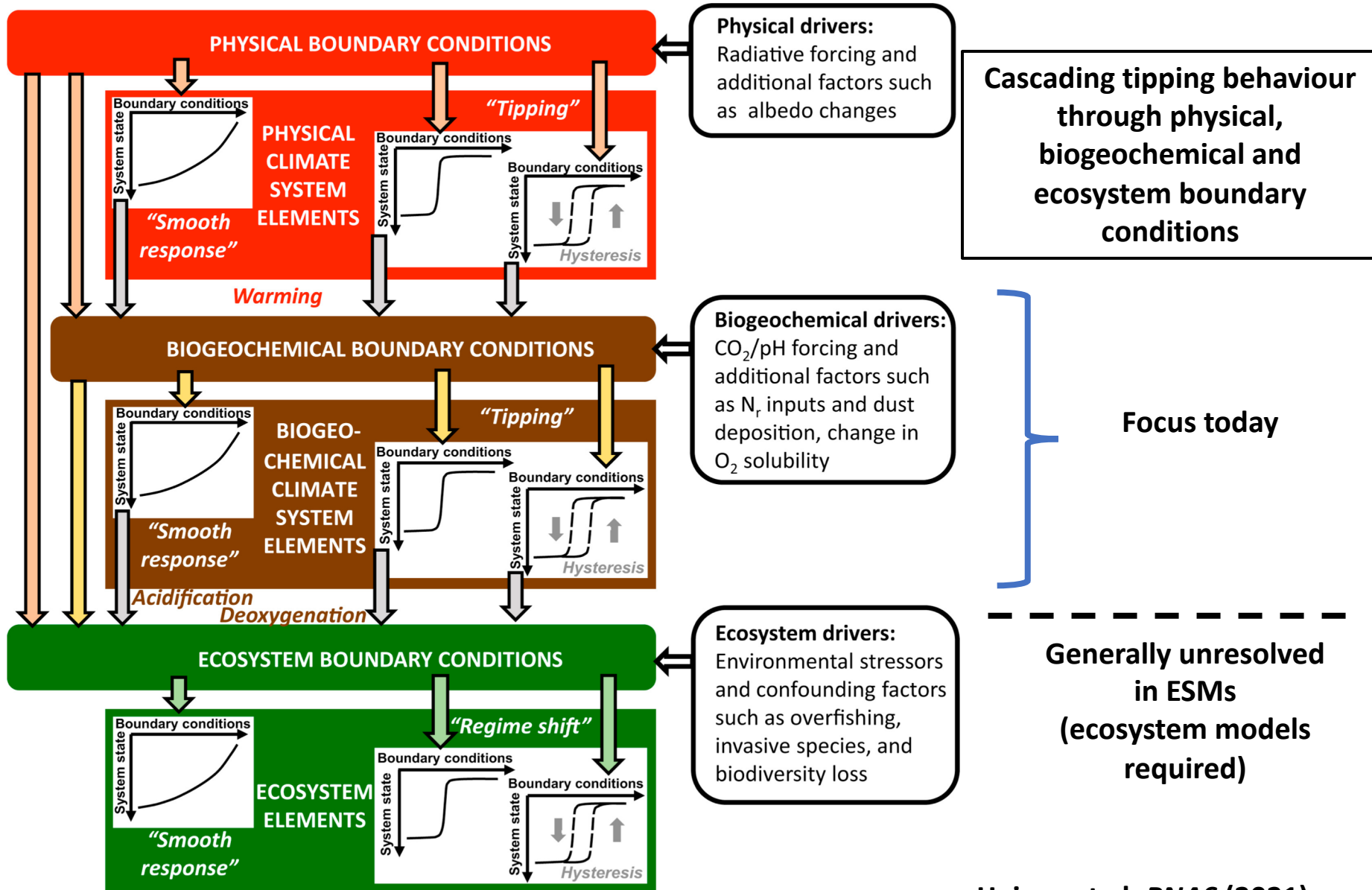


Institut  
Pierre  
Simon  
Laplace



Institut Henri Poincaré  
04/10/2023

# Ocean biogeochemical tipping points in ESMs: a regime shift in Arctic ocean chemistry ?



# IPCC SROCC identified MHWs, deoxygenation and acidification as potential abrupt ocean changes with variable reversibility

Change in system component	Potentially abrupt	Irreversibility if forcing reversed (time scales indicated)	Impacts on natural and human systems; global vs. regional vs. local	Projected likelihood and/or confidence level in 21st century under scenarios considered
<b>Ocean</b>				
Atlantic Meridional Overturning Circulation (AMOC) collapse (Section 6.7)	Yes	Unknown	Widespread; increased winter storms in Europe, reduced Sahelian rainfall and agricultural capacity, variations in tropical storms, increased sea levels on Atlantic coasts	<i>Very unlikely</i> , but physically plausible
Subpolar gyre (SPG) cooling (Section 6.7)	Yes	Irreversible within decades	Similar to AMOC impacts but considerably smaller	<i>Medium confidence</i>
Marine heatwave (MHW) increase (Section 6.4)	Yes	Reversible within decades to centuries	Coral bleaching, loss of biodiversity and ecosystem services, harmful algal blooms, species redistribution	<i>Very likely (very high confidence)</i> for physical change <i>High confidence</i> for impacts
Arctic sea ice retreat (Section 3.3)	Yes	Reversible within decades to centuries	Coastal erosion in Arctic (may take longer to reverse), impact on mid-latitude storms ( <i>low confidence</i> ); rise in Arctic surface temperatures ( <i>high confidence</i> )	<i>High confidence</i>
Ocean deoxygenation and hypoxic events (Section 5.2)	Yes	Reversible at surface, but irreversible for centuries to millennia at depth	Major changes in ocean productivity, biodiversity and biogeochemical cycles	<i>Medium confidence</i>
Ocean acidification (Section 5.2)	Yes	Reversible at surface, but irreversible for centuries to millennia at depth	Changes in growth, development, calcification, survival and abundance of species, for example, from algae to fish	<i>Virtually certain (very high confidence)</i>



**Tipping point: “A level of change in system properties beyond which a system reorganizes, often in a non-linear manner, and does not return to the initial state even if the drivers of the change are abated.”**  
(IPCC, SROCC, 2019)

# IPCC SROCC identified MHWs, deoxygenation and acidification as potential abrupt ocean changes with variable reversibility

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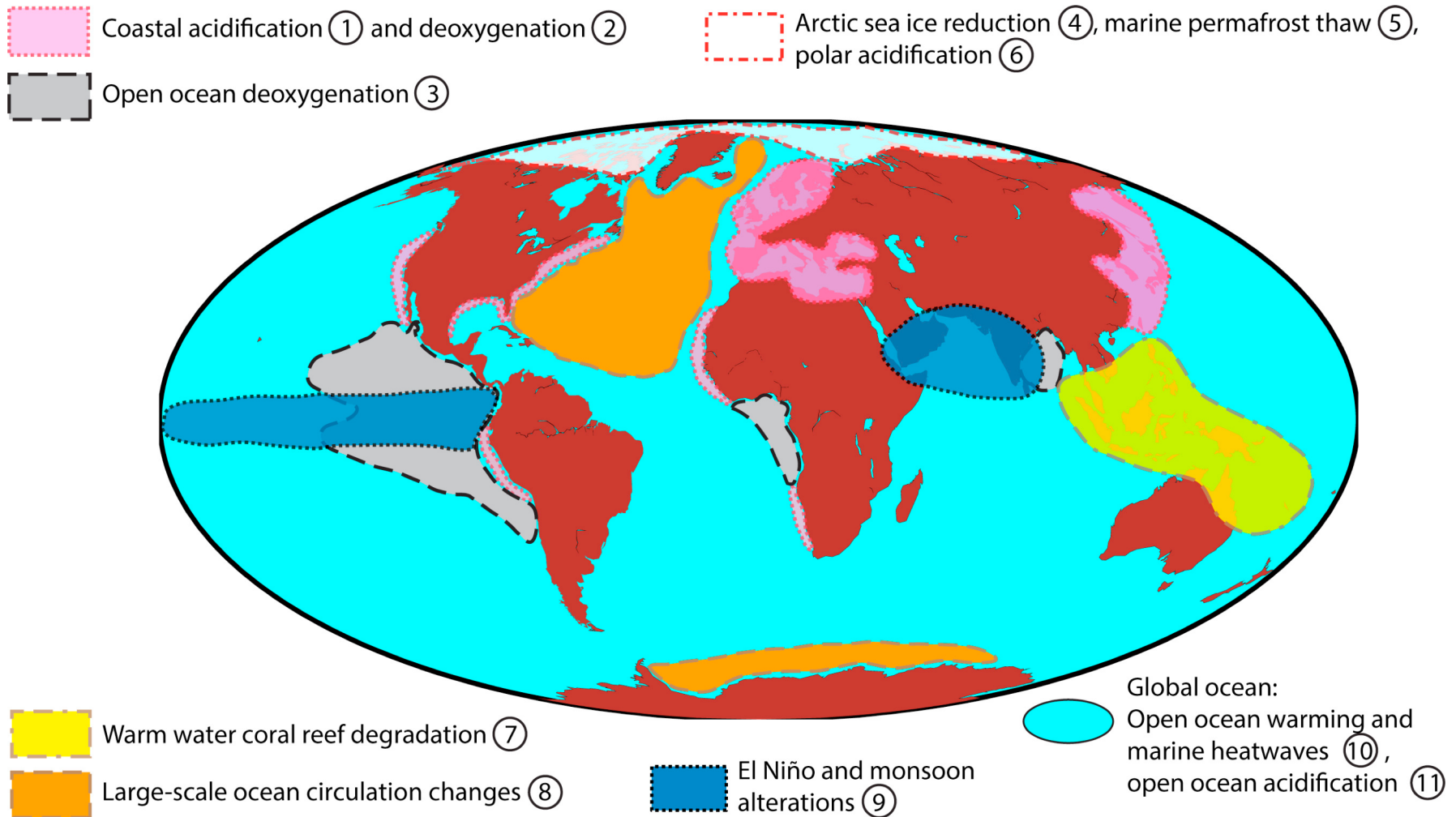
## Reversibility timescales

Marine heat waves -decades/centuries

Deoxygenation – reversible at surface, 10<sup>2</sup>-10<sup>3</sup> yrs at depth

Acidification – reversible at surface, 10<sup>2</sup>-10<sup>3</sup> yrs at depth

# Potential ocean tipping elements related to warming, deoxygenation and acidification



# Key biogeochemical abrupt change/tipping point papers to emerge from the COMFORT project



Our common future ocean in the Earth system – quantifying coupled cycles of carbon, oxygen, and nutrients for determining and achieving safe operating spaces with respect to tipping points: COMFORT (Horizon 2020)

SCIENCE ADVANCES | RESEARCH ARTICLE

ECOLOGY

## Abrupt shifts in 21st-century plankton communities

B. B. Cael<sup>1\*</sup>, Stephanie Dutkiewicz<sup>2</sup>, Stephanie Henson<sup>1</sup>

Research Article | Open Access |

## Biogeochemical Timescales of Climate Change Onset and Recovery in the North Atlantic Interior Under Rapid Atmospheric CO<sub>2</sub> Forcing

Leonardo Bertini , Jerry Tjiputra

Perspective

## Biogeochemical extremes and compound events in the ocean

<https://doi.org/10.1038/s41586-021-03981-7> Nicolas Gruber<sup>1,2</sup>, Philip W. Boyd<sup>3</sup>, Thomas L. Frölicher<sup>3,4</sup> & Meike Vogt<sup>1</sup>

Article

## Arctic Ocean annual high in $p_{\text{CO}_2}$ could shift from winter to summer

<https://doi.org/10.1038/s41586-022-05205-y> James C. Orr<sup>1,2,3</sup>, Lester Kwiatkowski<sup>2</sup> & Hans-Otto Pörtner<sup>3</sup>

LETTER • OPEN ACCESS

## Hysteresis of the Earth system under positive and negative CO<sub>2</sub> emissions

Aurich Jeltsch-Thömmes<sup>2</sup> , Thomas F. Stocker and Fortunat Joos

Published 3 December 2020 • © 2020 The Author(s). Published by IOP Publishing Ltd

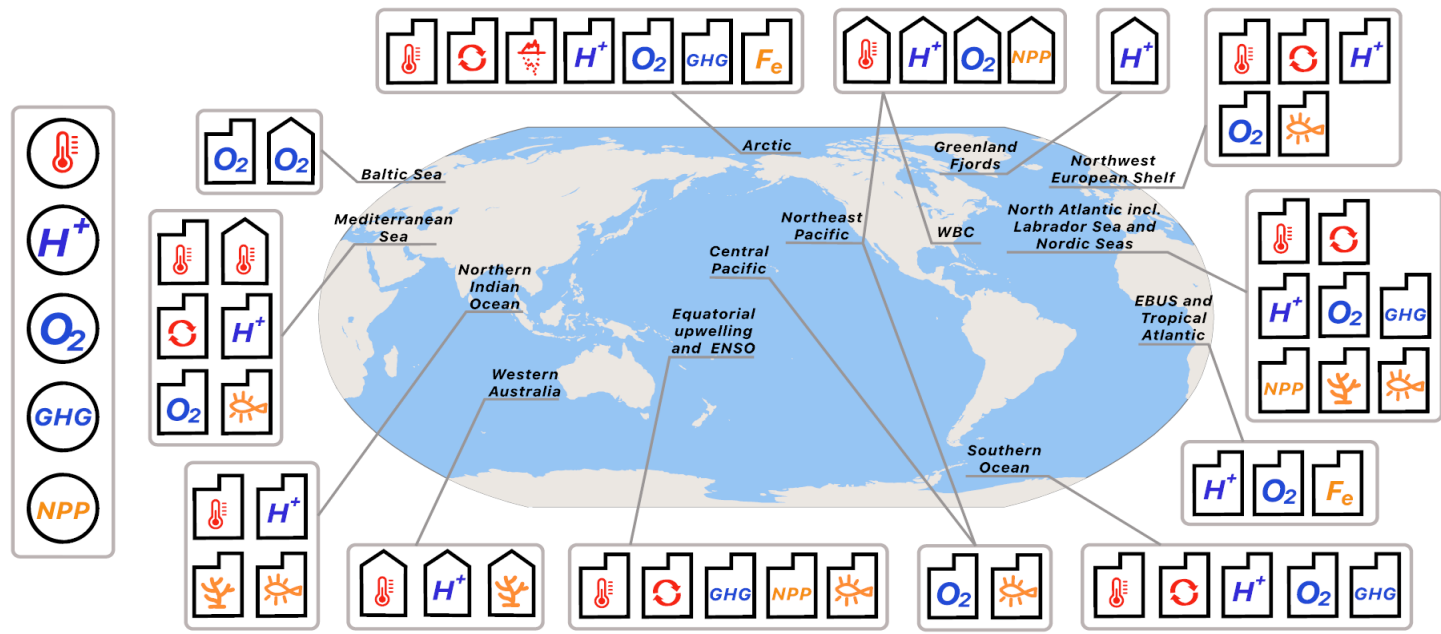
[Environmental Research Letters](#), Volume 15, Number 12

Extensive OBGC publications

But - **reversibility rarely addressed** and if so typically in EMICs not ESMs (indicative of ESM simulations performed in CMIP5/6?)  
- **early warning indicators receive limited attention**

# Key biogeochemical abrupt change/tipping point papers to emerge from the COMFORT project

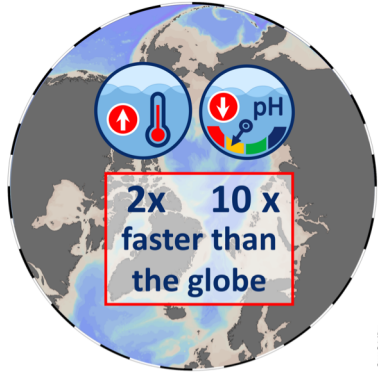
## Synthesis/review of ocean biogeochemical abrupt change/tipping points



- | Spatial/temporal behavior | Physical changes | Biogeochemical                   | Ecosystem related      |
|---------------------------|------------------|----------------------------------|------------------------|
| ○ global/millennial       | 🌡️ warming       | H <sup>+</sup> acidification     | NPP primary production |
| 📄 regional abrupt/decades | 🔄 circulation    | O <sub>2</sub> deoxygenation     | Fe Fe cycle            |
| 🏠 regional extremes       | 🌨️ sea ice       | GHG greenhouse gas sources/sinks | 🪸 corals               |
|                           |                  |                                  | 🌿 ecosystem health     |

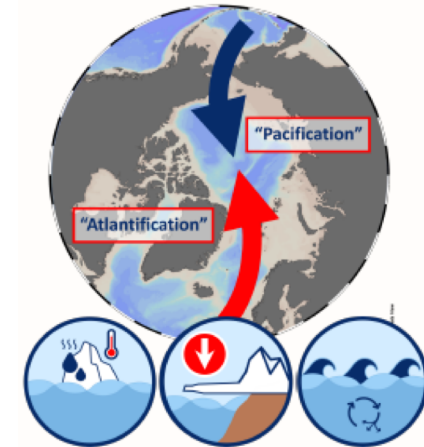


# COMFORT reiterated the sensitivity of the Arctic Ocean to abrupt change and potential tipping points



**Warming 2x faster and acidifying  
10x faster than the global mean**

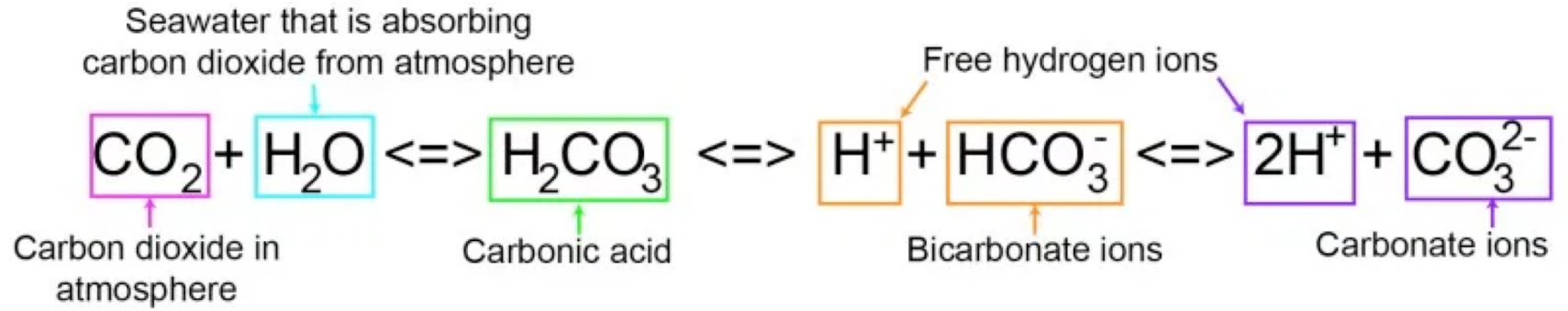
**Arctic Ocean currents are changing due to increased  
freshwater inputs and are invaded by the waters from  
the Atlantic and Pacific Oceans**



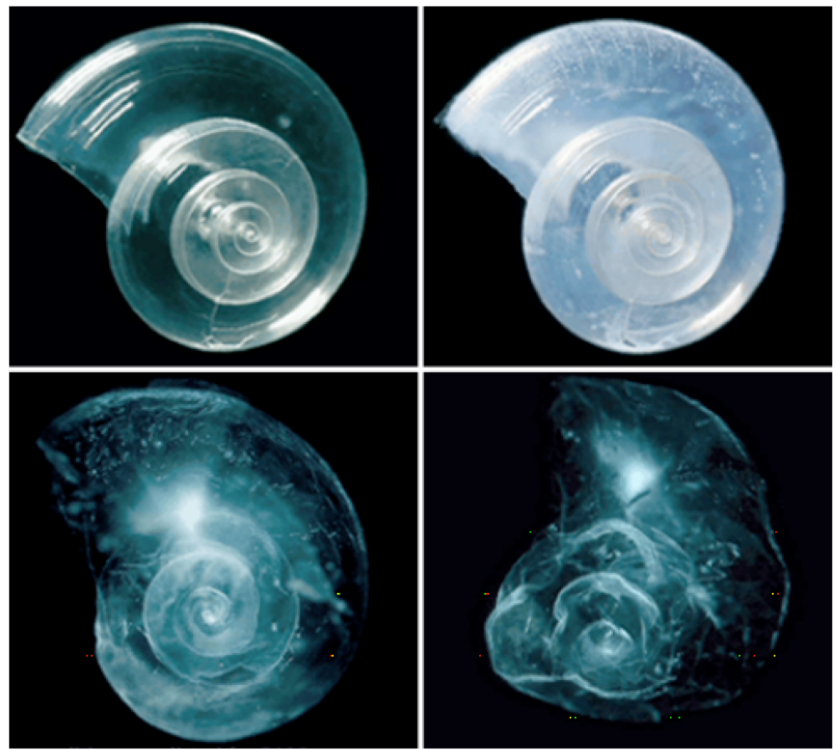
**Extreme events expected to become more frequent  
and intense, due to summer sea-ice & Greenland ice  
sheet loss, wildfires, and permafrost thaw**



# Ocean acidification and why we might care about it



## Pteropod dissolving in $\Omega_{\text{arag}} < 1$ seawater



- $p\text{CO}_2$  ↑
- $\text{HCO}_3^-$  ↑
- $\text{H}^+$  ↑
- $\text{CO}_3^{2-}$  ↓
- pH ↓
- $\Omega_{\text{arag}}$  ↓
- $\Omega_{\text{calc}}$  ↓

# Temporal variability: Seasonal phasing of $p\text{CO}_2$ in the Arctic Ocean

**Ocean  $p\text{CO}_2$  is increasing**

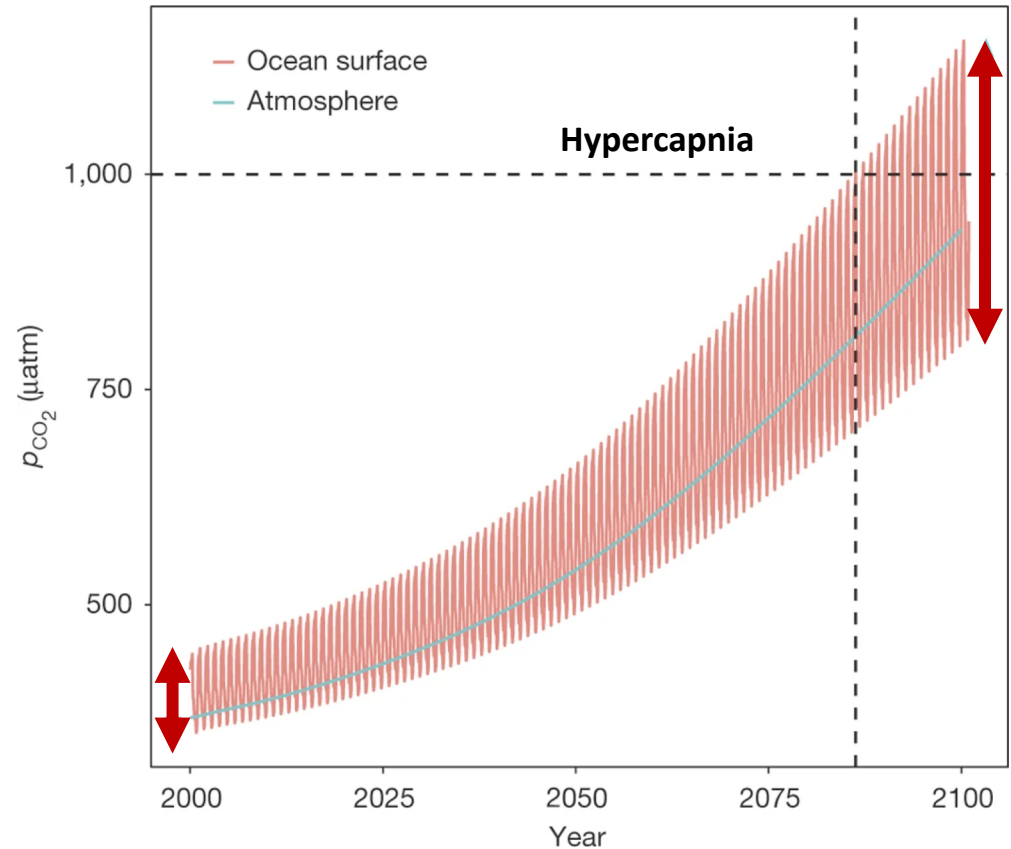
- focus here is on seasonal cycle

**Changing seasonal cycles can...**

- Influence when critical organism thresholds are crossed e.g.

- hypercapnia ( $p\text{CO}_2$ )
- Acidosis ( $\text{H}^+$ )
- $\Omega_{\text{arag}}$  undersaturation

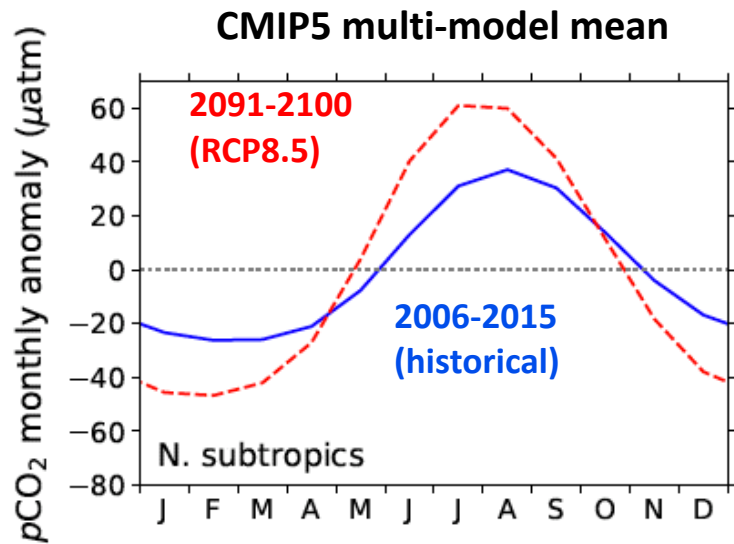
- Impact air-sea gas exchange



# Amplification of $p\text{CO}_2$ seasonal cycle in the surface ocean

The  $p\text{CO}_2$  seasonal cycle is projected to amplify in the surface ocean

...but seasonal phasing (max/mins) near identical



Amplification supported by observations (e.g. Landschutser et al., NCC, 2018)

typically thermally driven seasonal cycles of  $p\text{CO}_2$  in oligotrophic regions

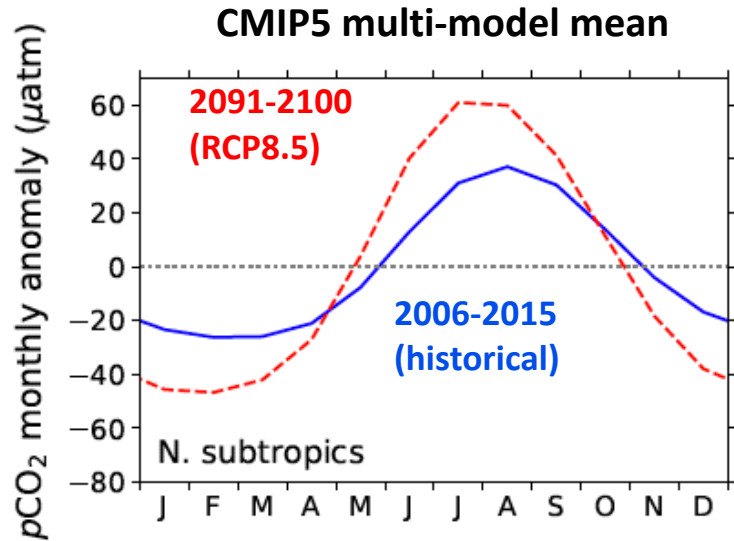
Kwiatkowski & Orr, *NCC* (2018)

Orr, Kwiatkowski & Pörtner, *Nature* (2022)

# Amplification of $p\text{CO}_2$ seasonal cycle in the surface ocean

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...but seasonal phasing (max/mins) near identical



Amplification driven by ocean carbon uptake (not climate change)

**i.e. Geochemical not Radiative**

-seen in “ $\text{CO}_2$ -only” simulations where model radiative module see constant  $\text{CO}_2$  (e.g. esmfixclim1)

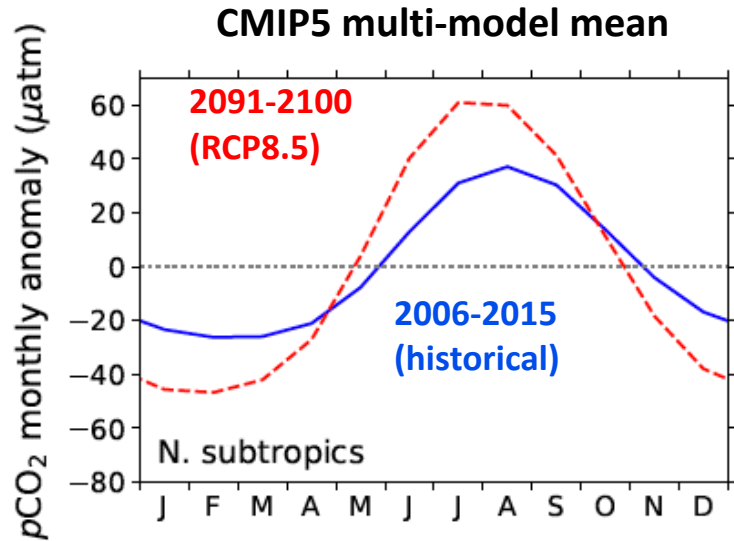
Kwiatkowski & Orr, *NCC* (2018)

Orr, Kwiatkowski & Pörtner, *Nature* (2022)

# Amplification of $p\text{CO}_2$ seasonal cycle in the global surface ocean

The  $p\text{CO}_2$  seasonal cycle is projected to amplify in the surface ocean

...but seasonal phasing (max/mins) near identical



Amplification driven by ocean carbon uptake (not climate change)

i.e. Geochemical not Radiative

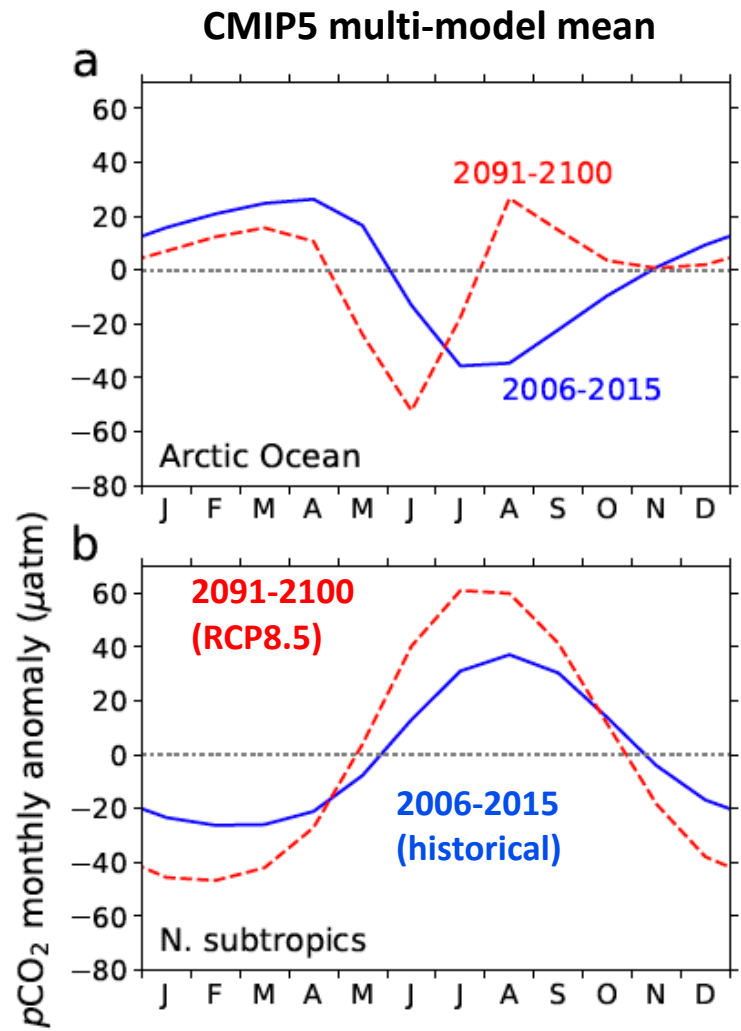
$$p'_{\text{CO}_2} \approx \frac{\partial p_{\text{CO}_2}}{\partial T} T' + \frac{\partial p_{\text{CO}_2}}{\partial S} S' + \frac{\partial p_{\text{CO}_2}}{\partial A_{\text{T}}} A'_{\text{T}} + \frac{\partial p_{\text{CO}_2}}{\partial C_{\text{T}}} C'_{\text{T}}$$

Ocean carbon uptake increases the sensitivity of  $p\text{CO}_2$  to it's driving variables

Kwiatkowski & Orr, *NCC* (2018)

Orr, Kwiatkowski & Pörtner, *Nature* (2022)

# But $p\text{CO}_2$ seasonal cycle “breakdown” in the Arctic Ocean

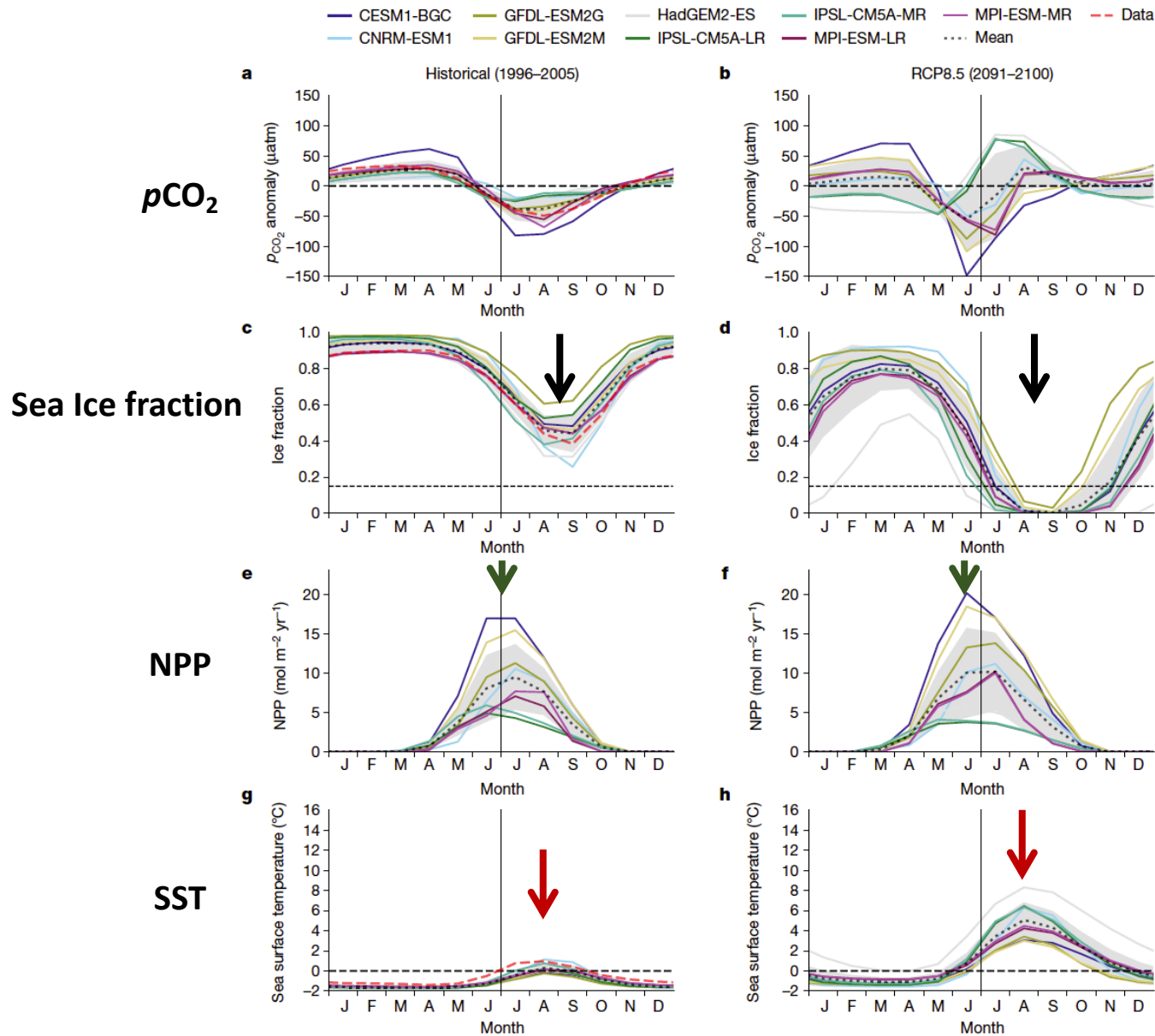


In the Arctic Ocean,

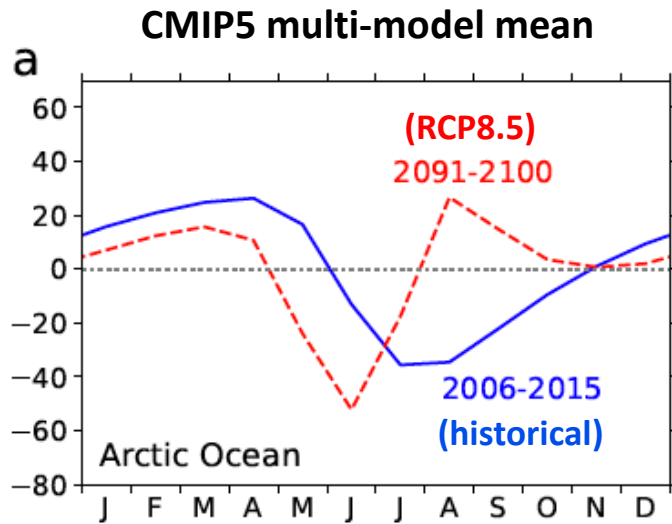
$p\text{CO}_2$  seasonal cycle is near-inverted  
Summer min becomes a summer max

$p\text{CO}_2$  summers are becoming winters  
Potential impact on sensitive species as  
this is the Arctic growth period

# Similar seasonal phasing of Arctic sea ice, NPP and SST (unlike $p\text{CO}_2$ ) but very different amplitudes



# Temporal variability: Seasonal phasing of $p\text{CO}_2$ in the Arctic Ocean



## Under RCP8.5

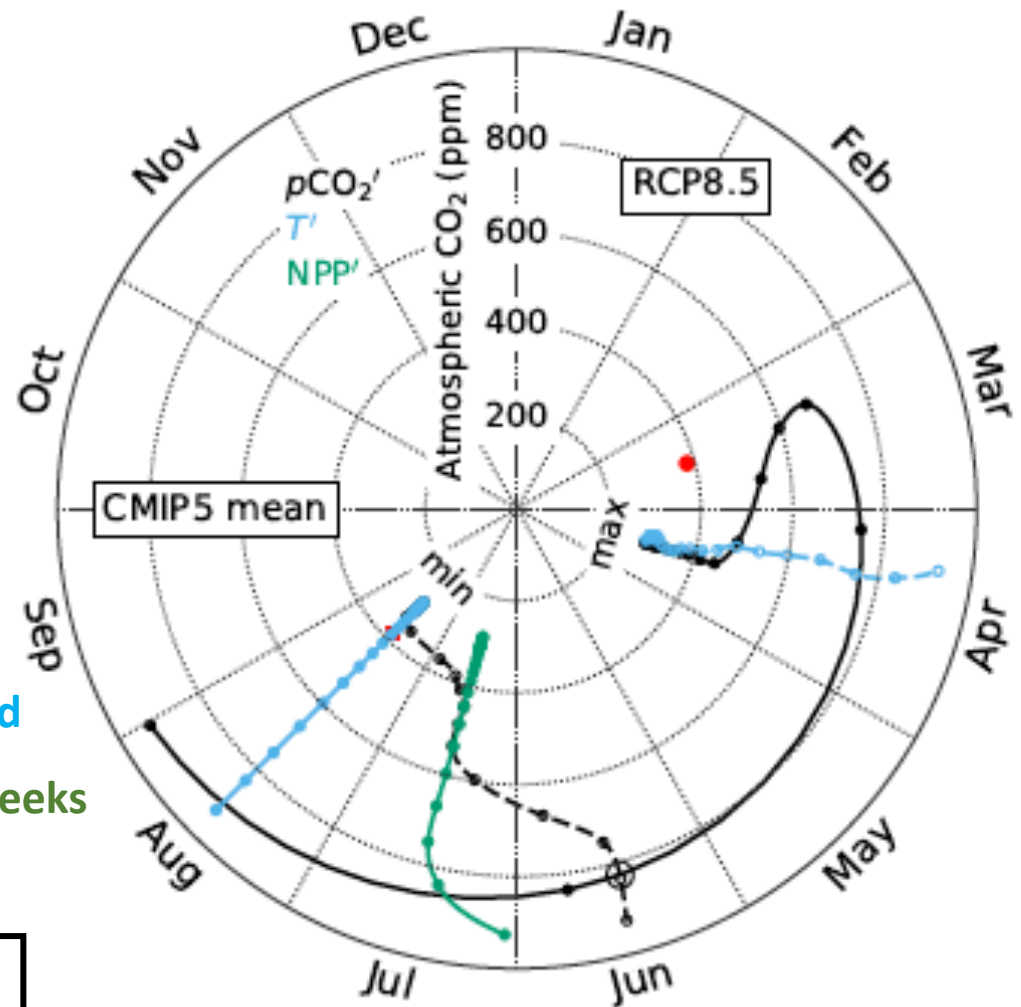
SST min and max months unchanged

Phytoplankton NPP max occurs 2-weeks  
(earlier spring bloom phenology)

However...

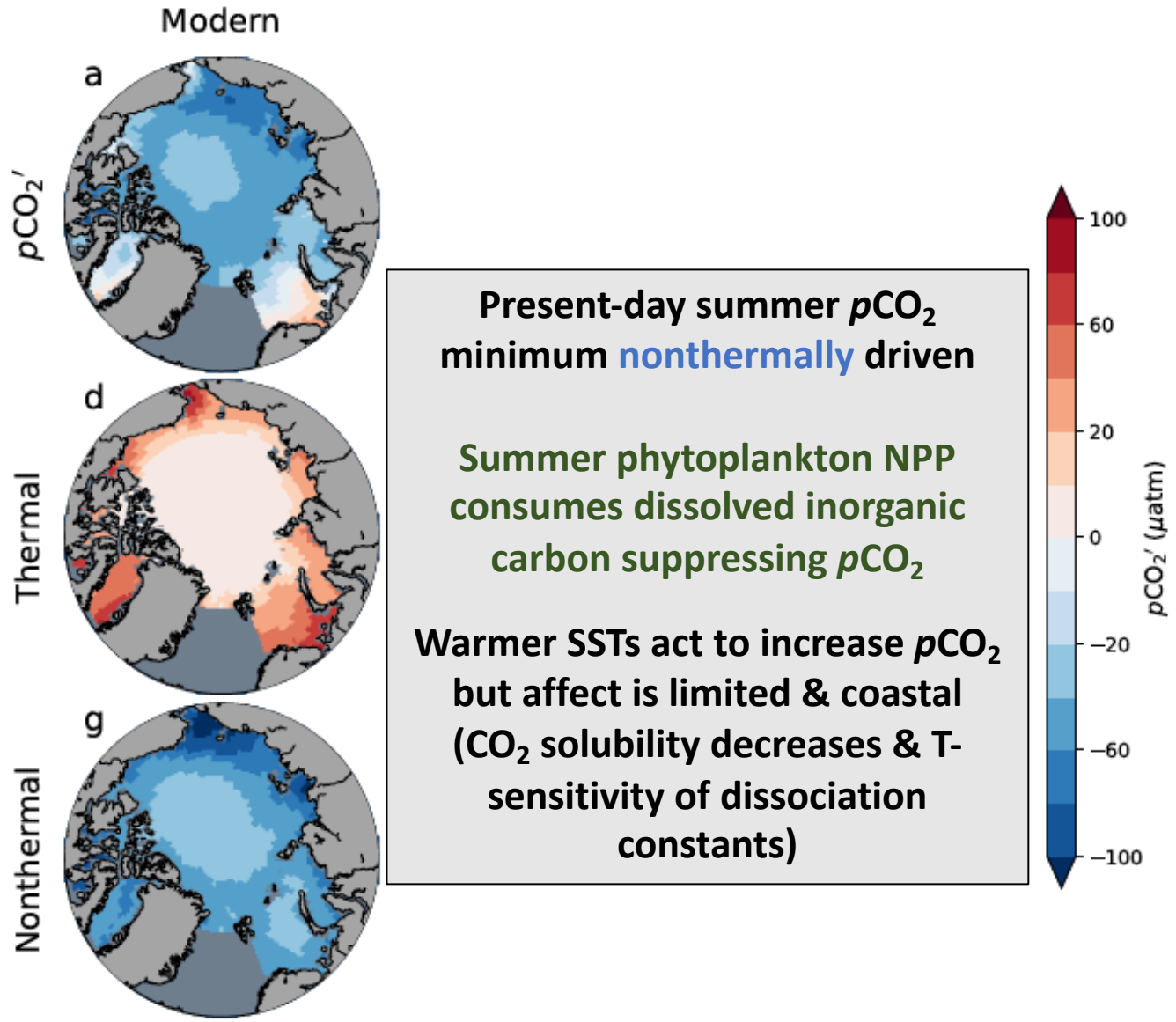
$p\text{CO}_2$  max occurs 4-months earlier

$p\text{CO}_2$  min occurs 2-months later





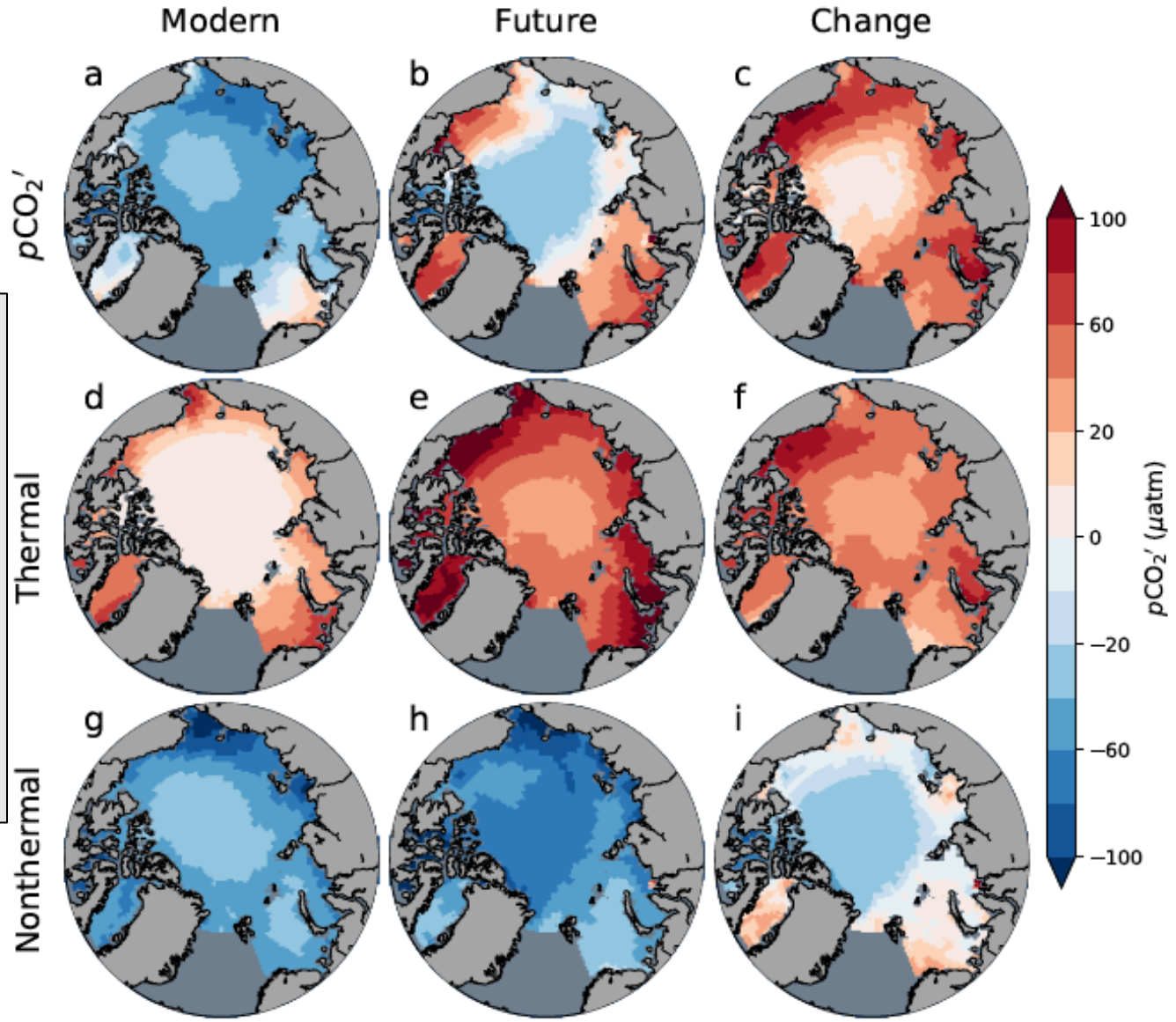
# Current summer seasonal $p\text{CO}_2$ minimum biologically driven



CMIP5 mean

# But future summer seasonal $p\text{CO}_2$ max is thermally driven

But in future (2091-2100) under high emissions,  
**Growth in thermal component overwhelms nonthermal**  
- particularly in coastal ocean



CMIP5 mean

# But future summer seasonal $p\text{CO}_2$ max is thermally driven

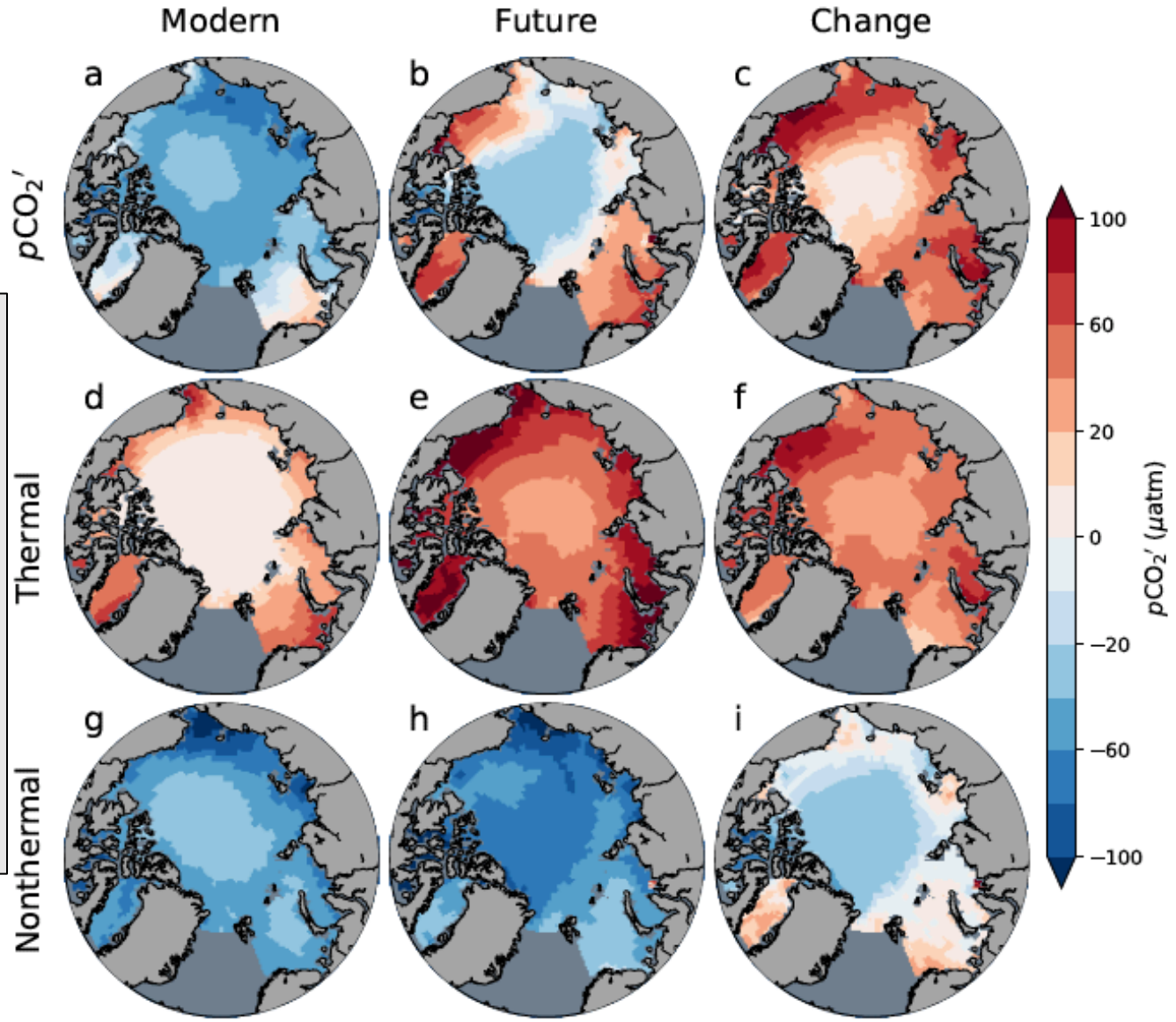
$$\left(\frac{\partial p\text{CO}_2}{\partial T}\right) \Delta T$$

↑

**Growth in thermal component is driven by seasonal SST anomalies (not thermal sensitivity)**

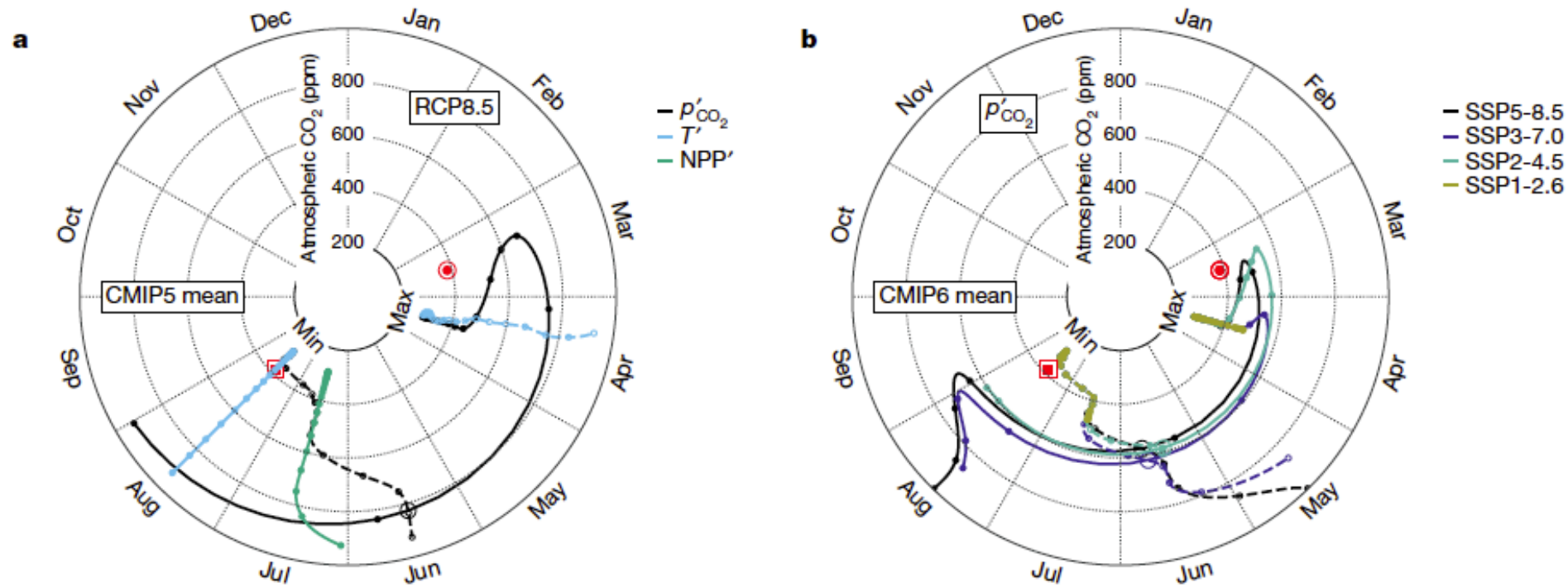
**i.e. climate change not ocean carbon uptake**

**(unlike global amplification)**



CMIP5 mean

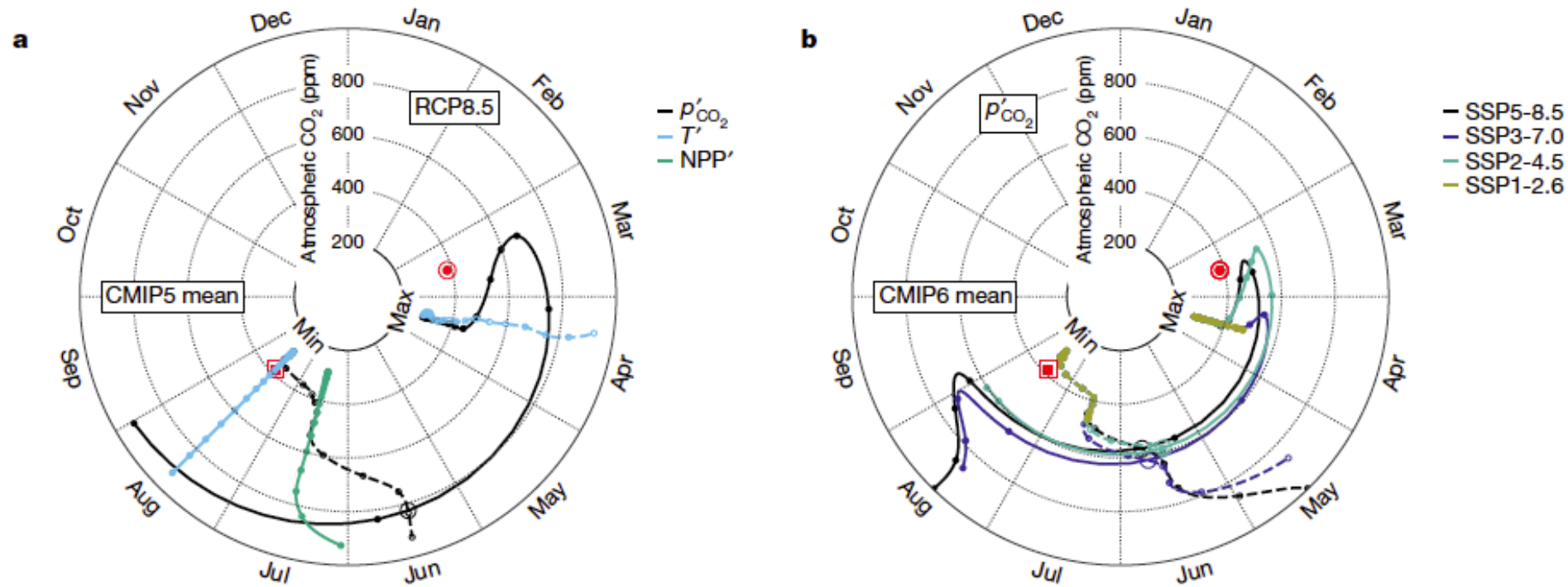
# Similar $p\text{CO}_2$ phase shift in CMIP6 (at lower atmospheric $\text{CO}_2$ )



Phase shift occurs at 500-600ppm in CMIP6  
(compared to 700-800ppm in CMIP5)

- Models have
- Higher ECS
  - Greater warming
  - Greater Arctic ice loss

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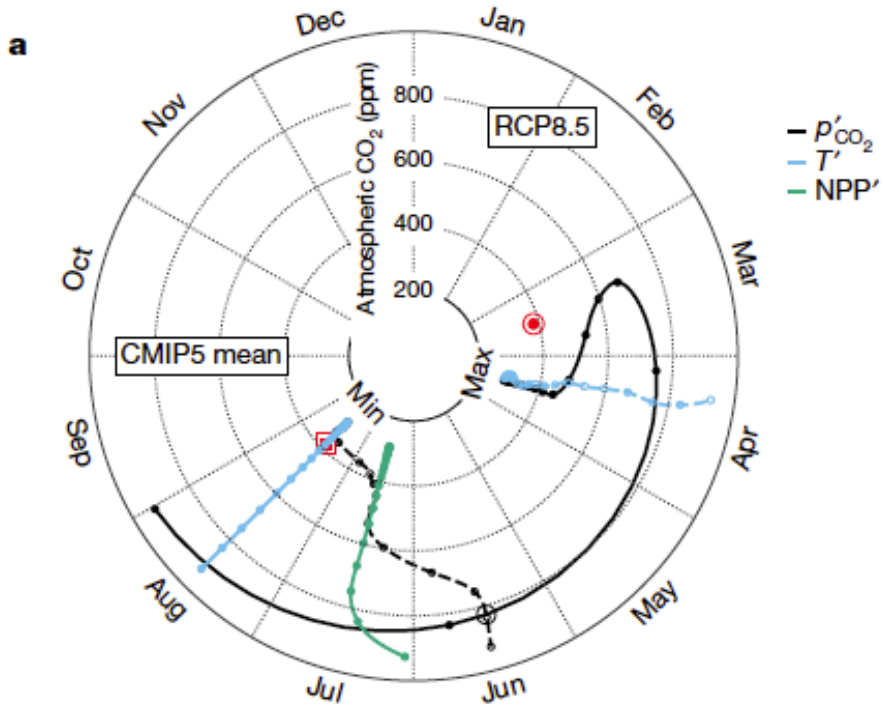
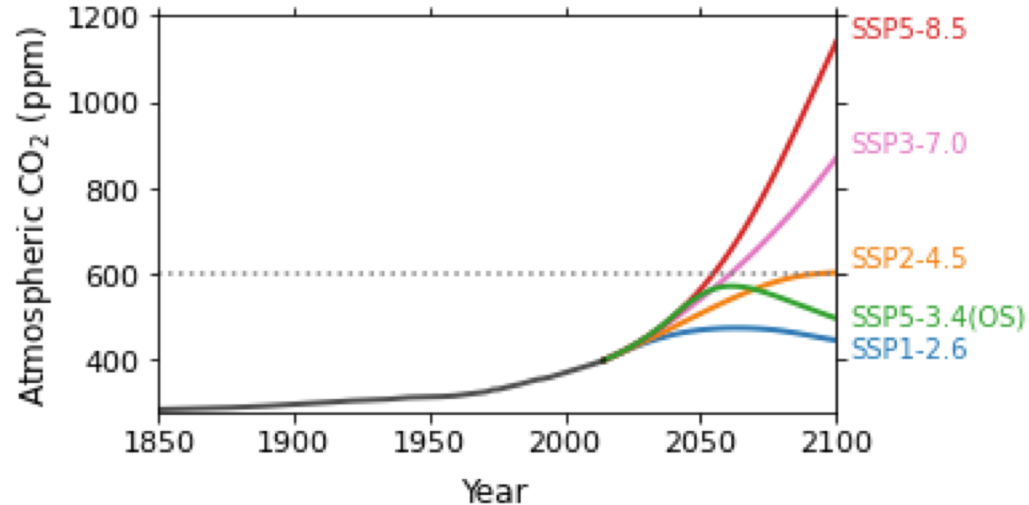
- Higher ECS
- Greater warming
- Greater Arctic ice loss



An opportunity to explore  
reversibility in overshoot  
scenarios  
(SSP5-3.4)

# Is the change in $p\text{CO}_2$ seasonality a tipping point/reversible?

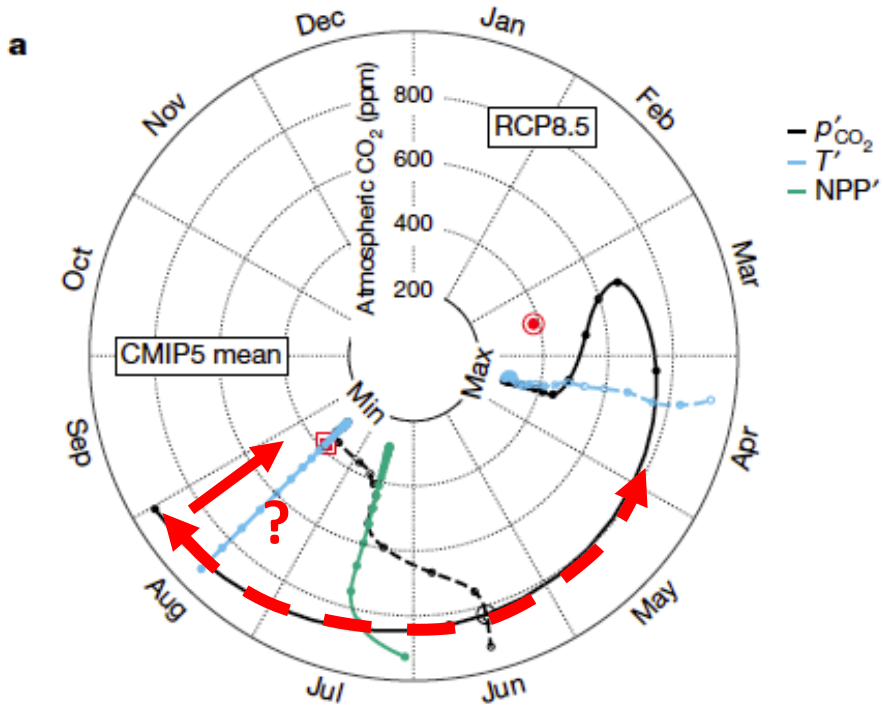
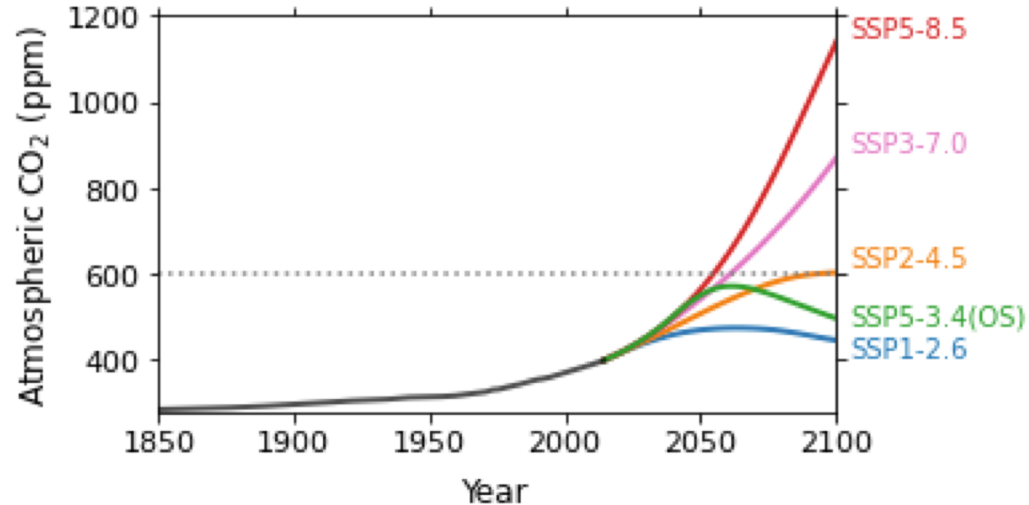
If atmospheric  $\text{CO}_2$  concentration is reduced does the seasonal cycle return to "normal" and how fast?



SSP5-3.4(OS) – peak  $\text{CO}_2$  570 ppm 2061 then decreases to 500 ppm 2100

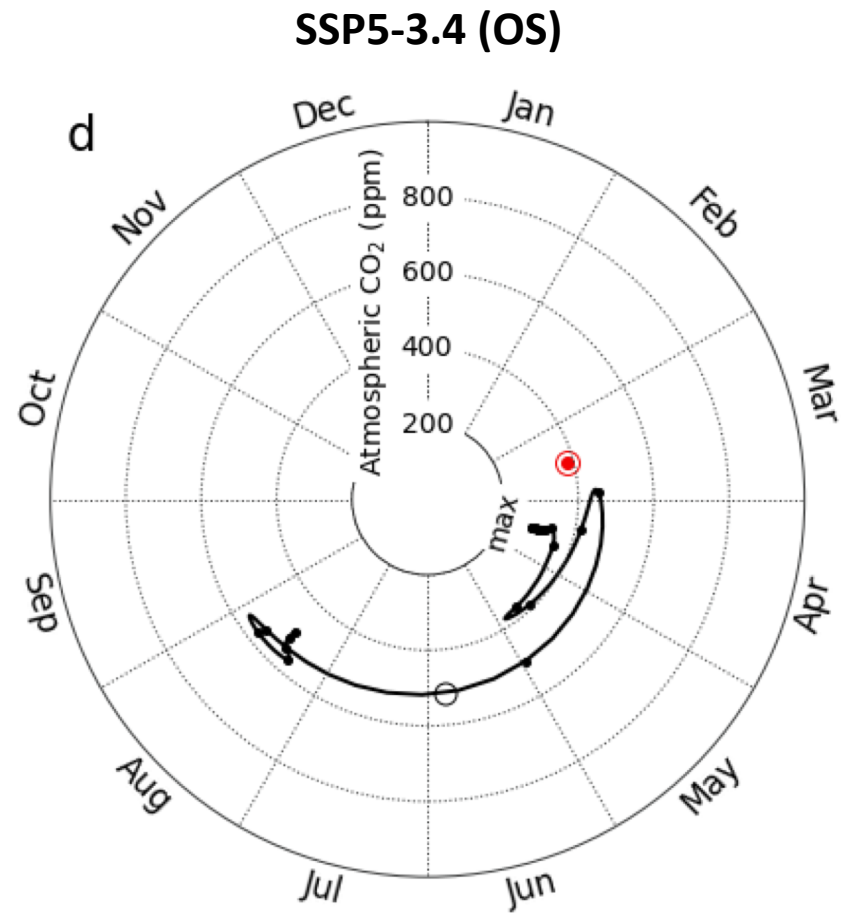
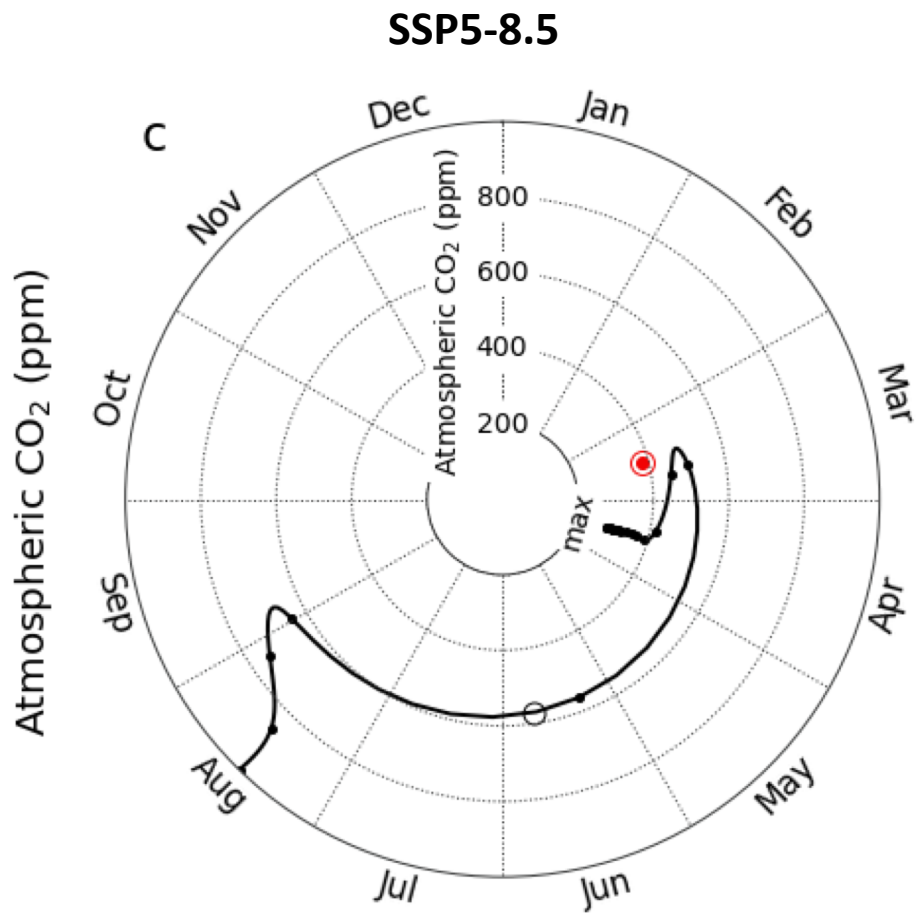
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SSP5-3.4(OS) – peak CO<sub>2</sub> 570 ppm 2061 then decreases to 500 ppm 2100

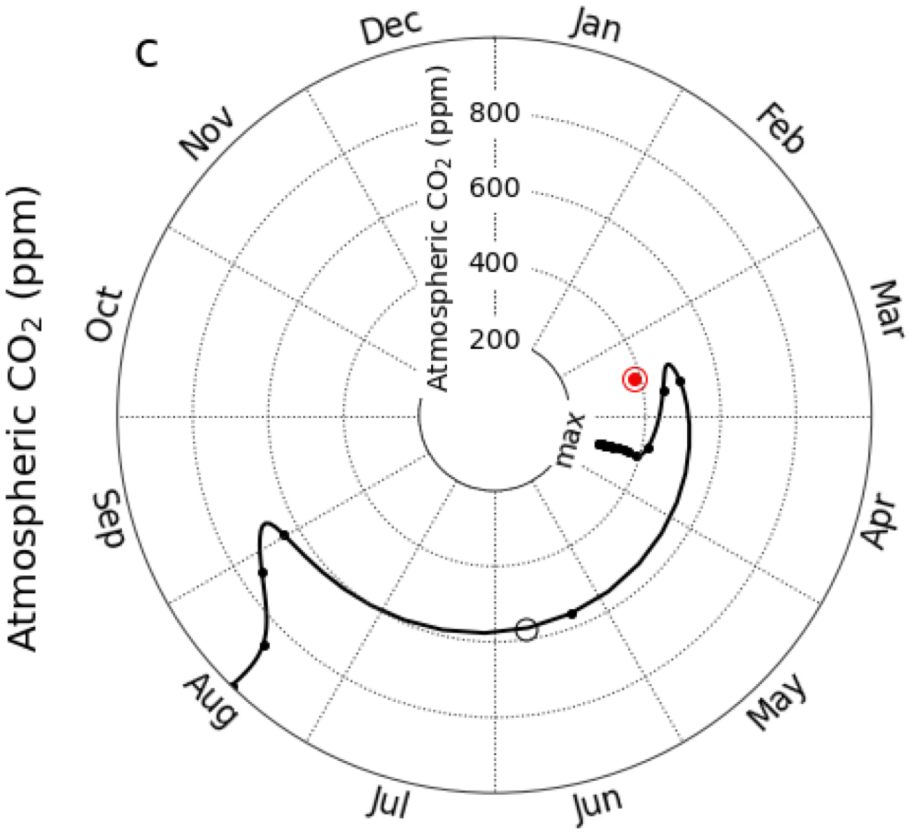
# Limited indication of reversibility in the multi-model mean



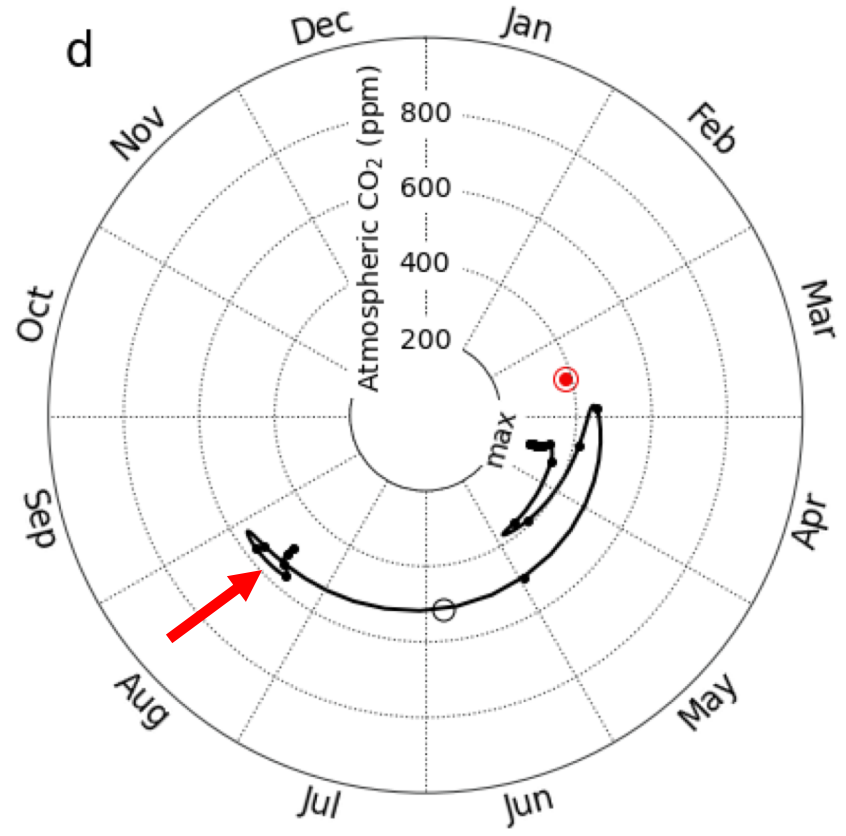


# Limited indication of reversibility in the multi-model mean

SSP5-8.5

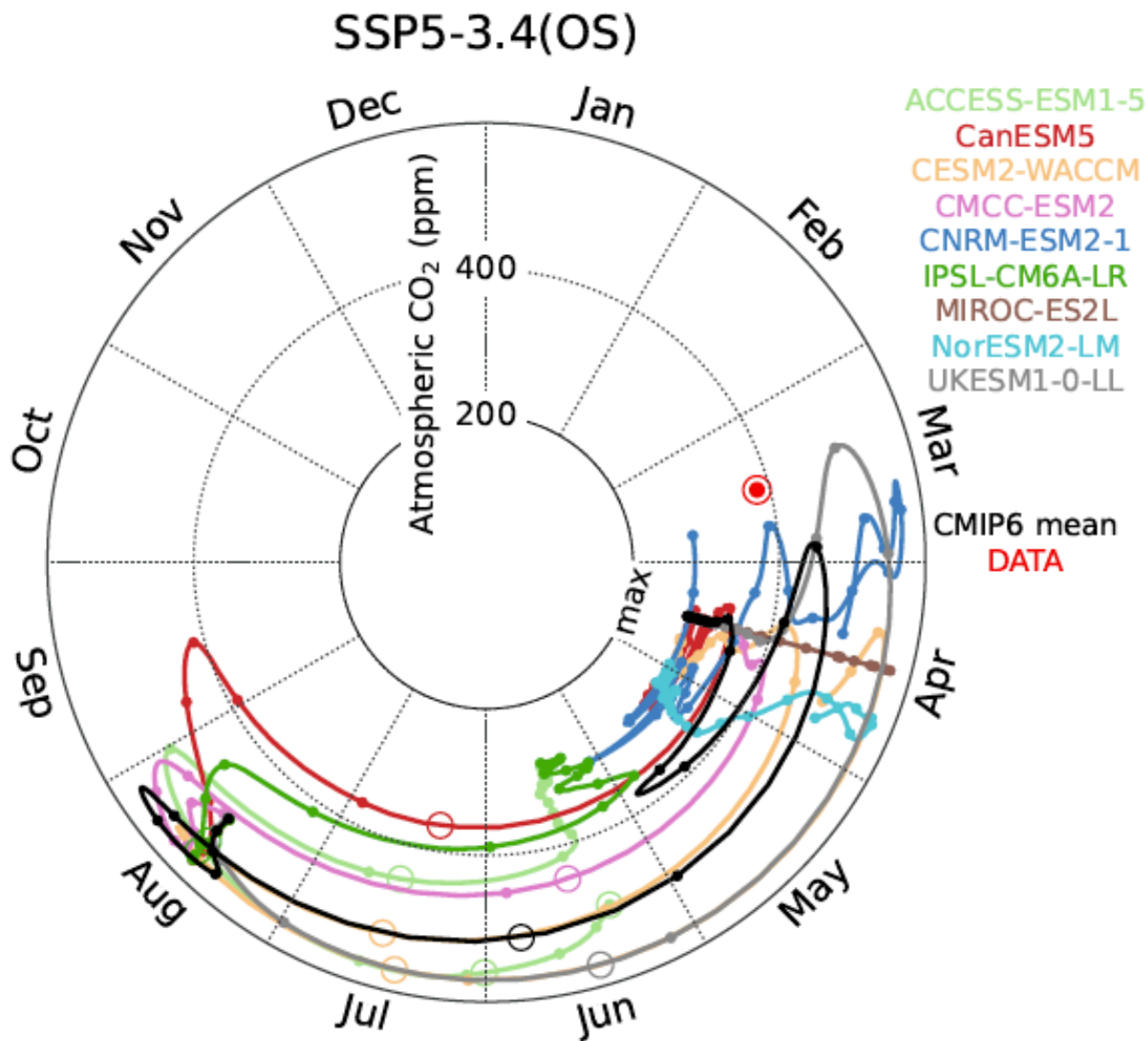


SSP5-3.4 (OS)

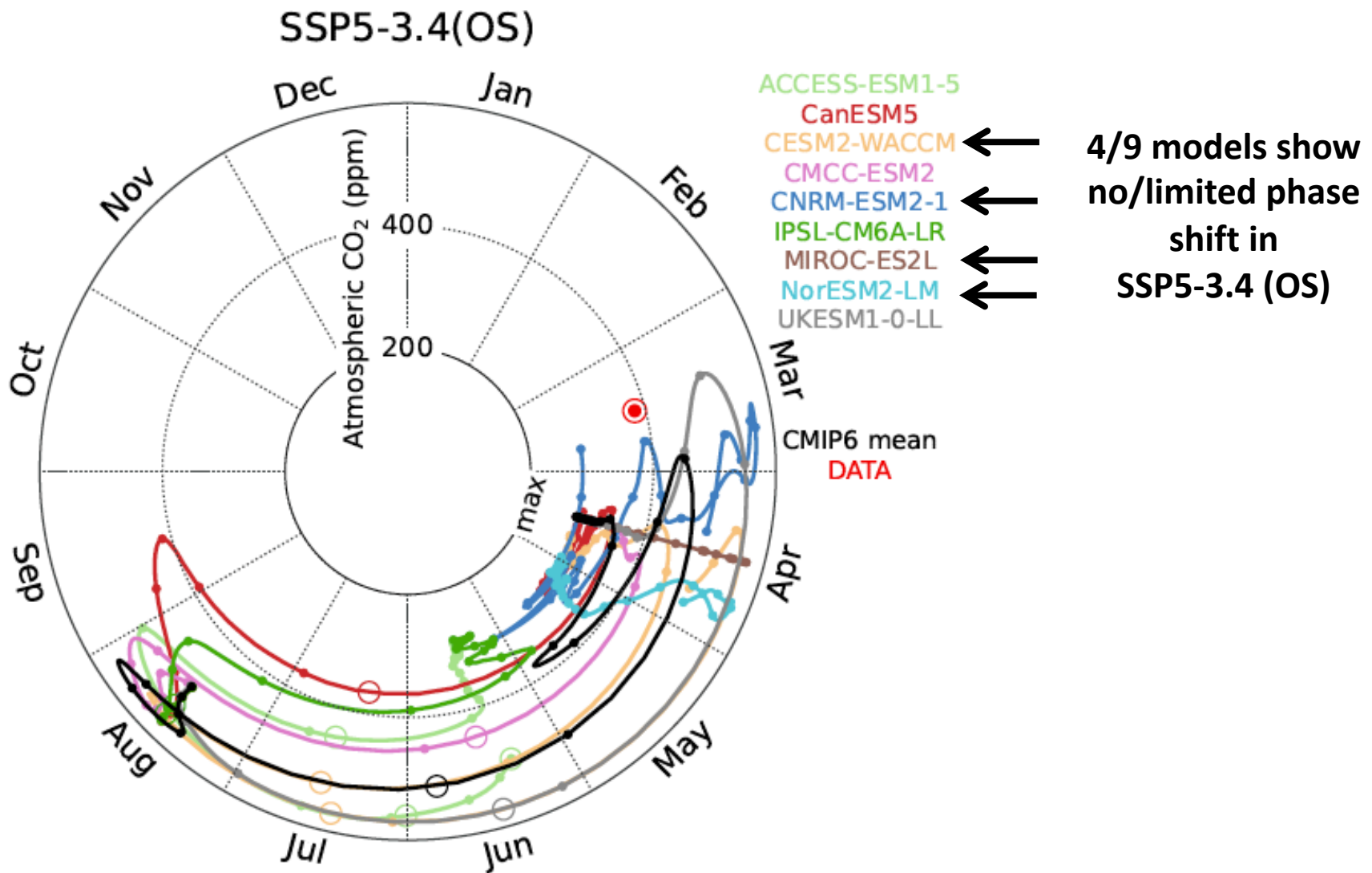


**$p\text{CO}_2$  max stays in August from 570 to 500 ppm  
(was in May previously at 500 ppm)  
Negative hysteresis? Irreversibility?**

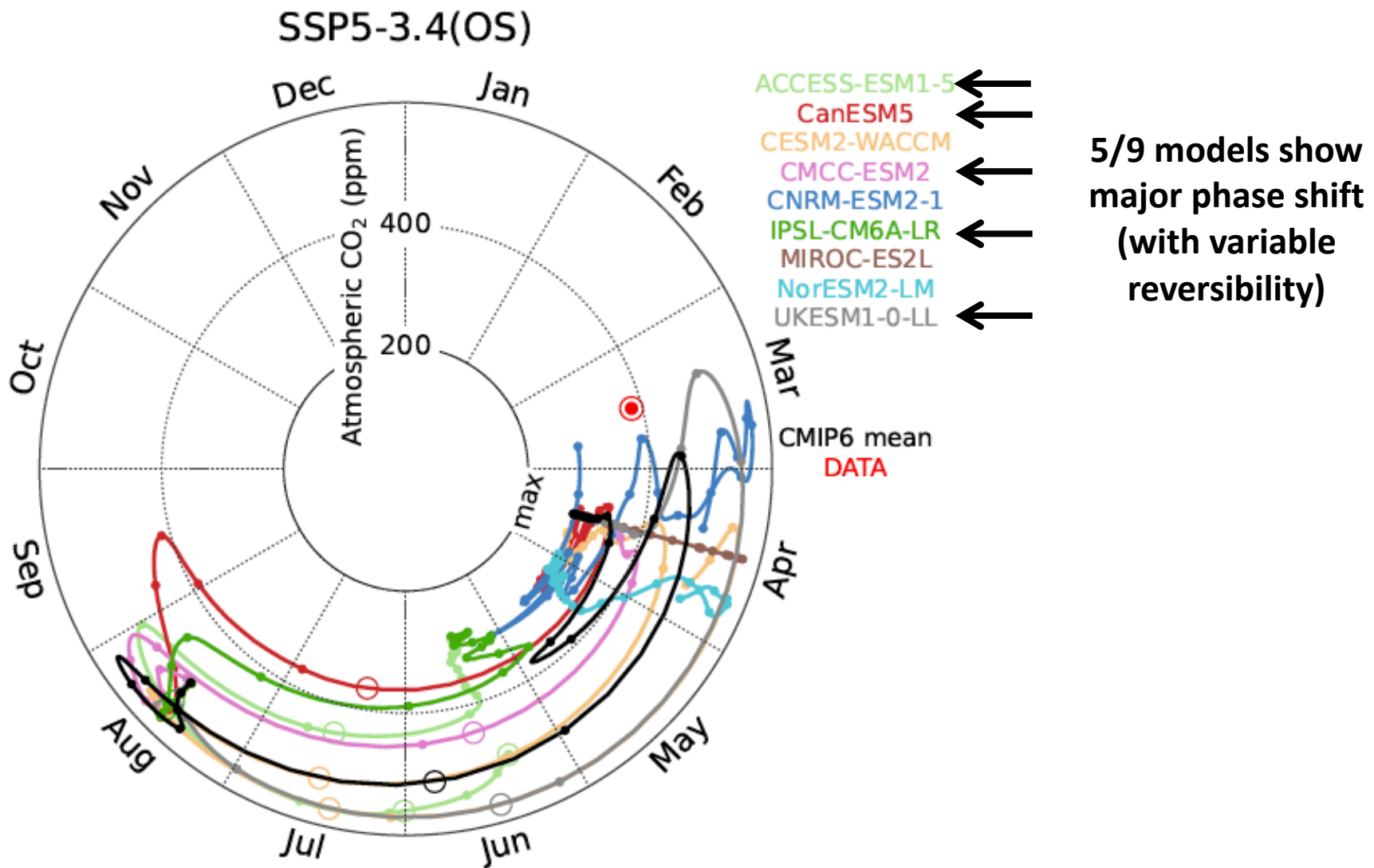
# Limited indication of reversibility in the multi-model mean: but lots of model uncertainty



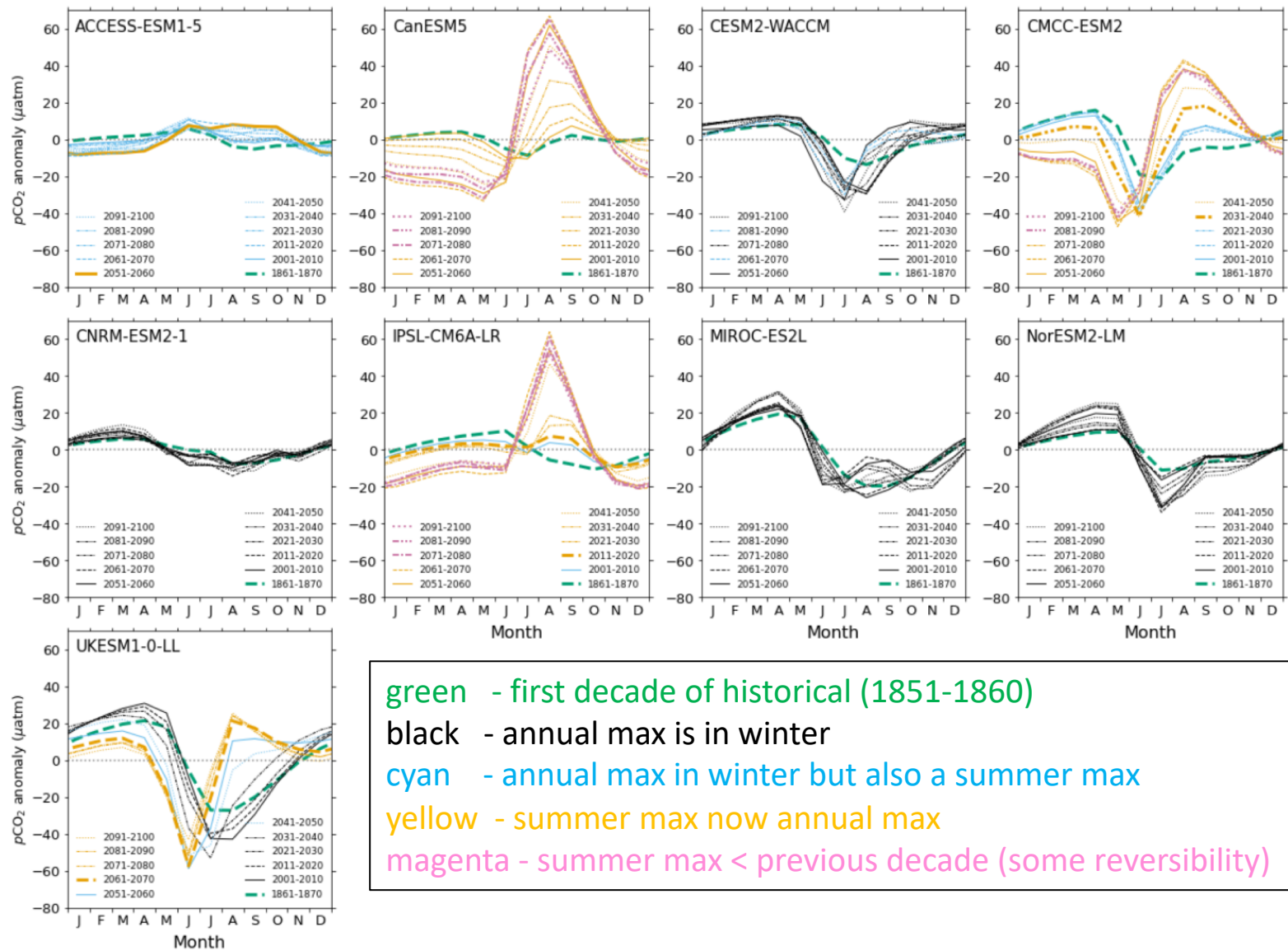
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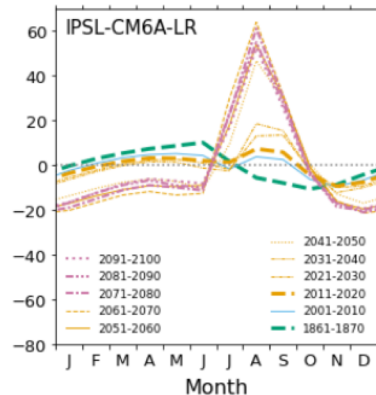
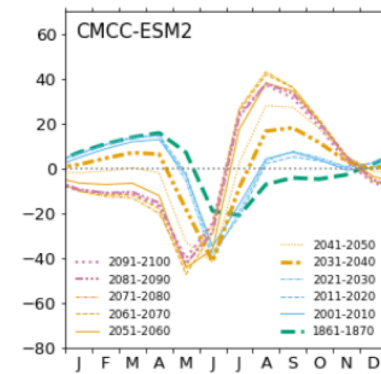
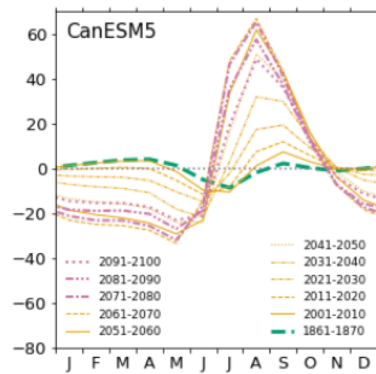
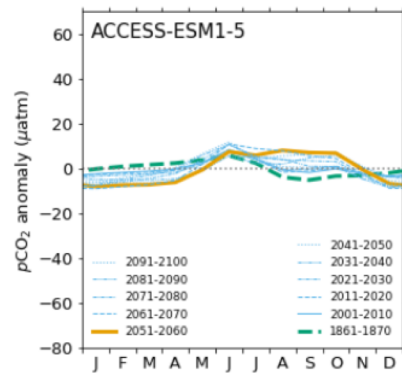
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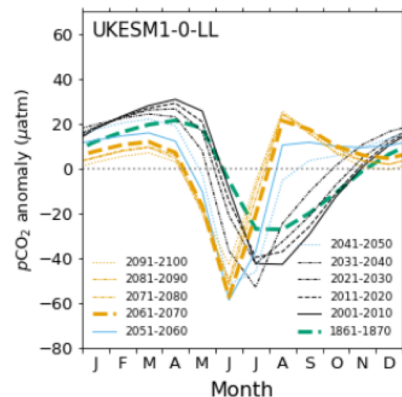
# Variable $p\text{CO}_2$ phase shift reversibility across CMIP6 models



# Variable $p\text{CO}_2$ phase shift reversibility across CMIP6 models

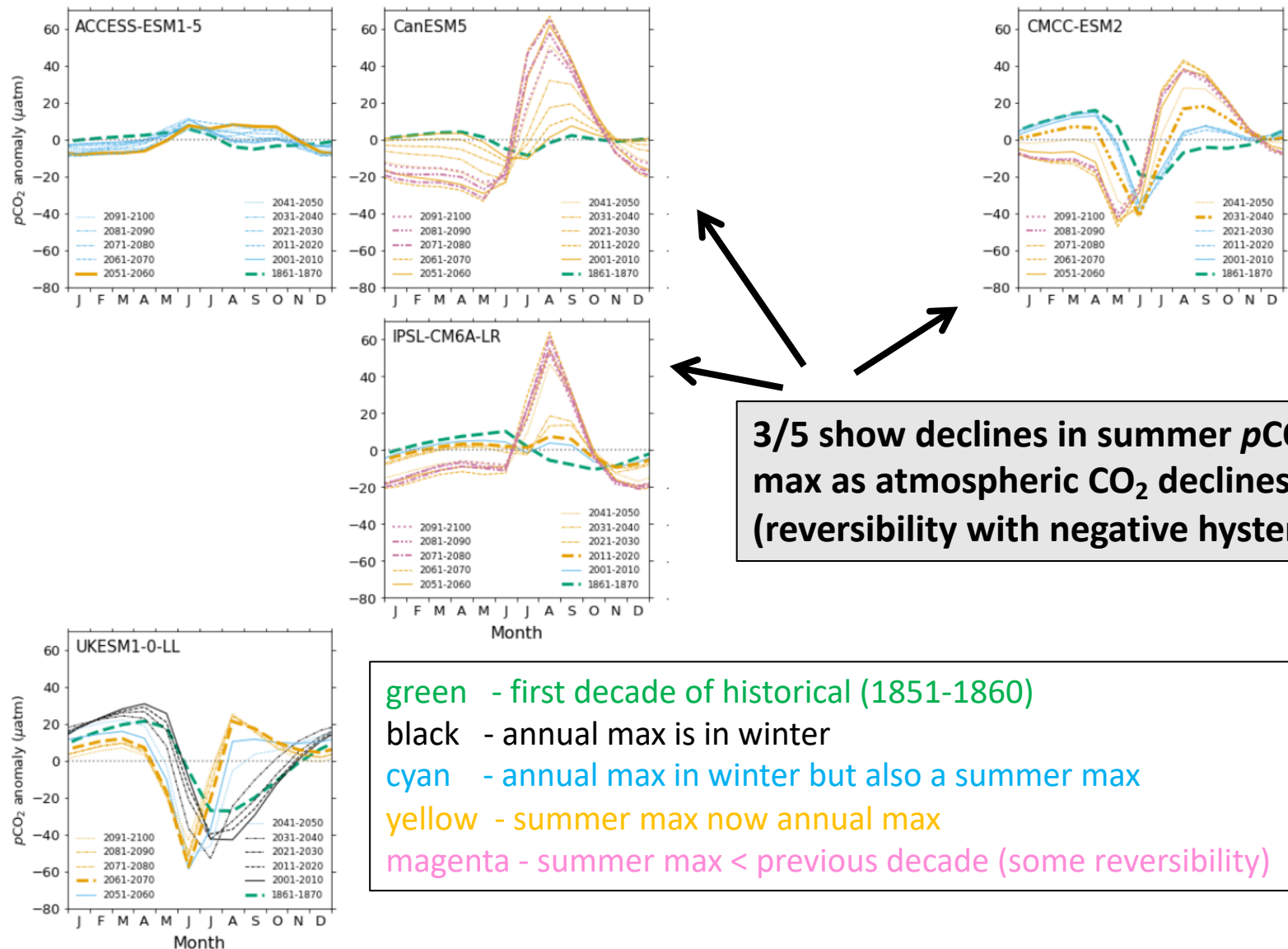


**5 models where the summer max becomes annual max**



green - first decade of historical (1851-1860)  
black - annual max is in winter  
cyan - annual max in winter but also a summer max  
yellow - summer max now annual max  
magenta - summer max < previous decade (some reversibility)

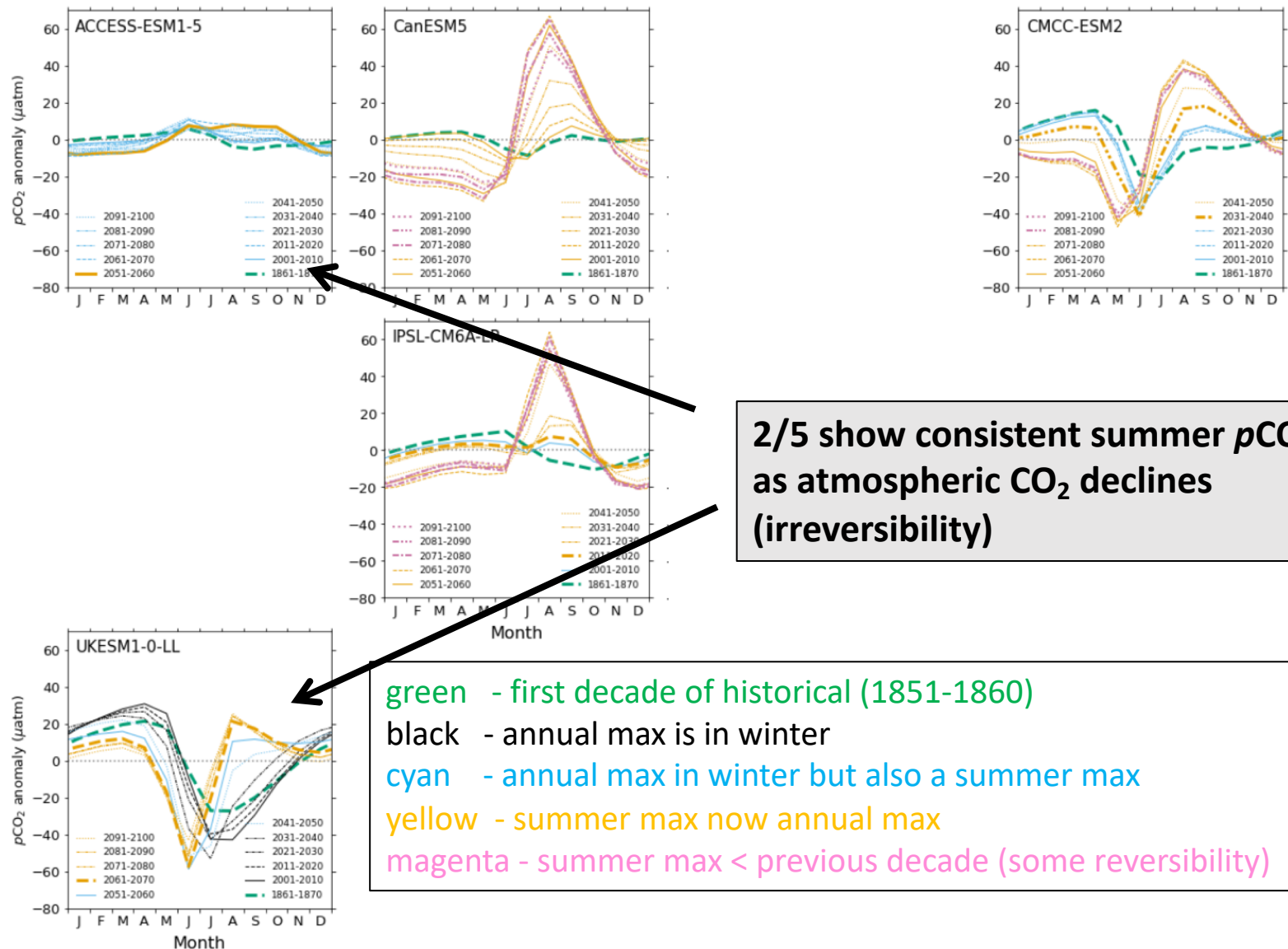
# Variable $p\text{CO}_2$ phase shift reversibility across CMIP6 models



**3/5 show declines in summer  $p\text{CO}_2$  max as atmospheric  $\text{CO}_2$  declines (reversibility with negative hysteresis?)**

green - first decade of historical (1851-1860)  
 black - annual max is in winter  
 cyan - annual max in winter but also a summer max  
 yellow - summer max now annual max  
 magenta - summer max < previous decade (some reversibility)

# Variable $p\text{CO}_2$ phase shift reversibility across CMIP6 models

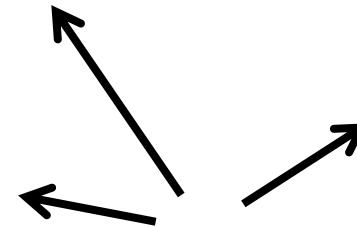
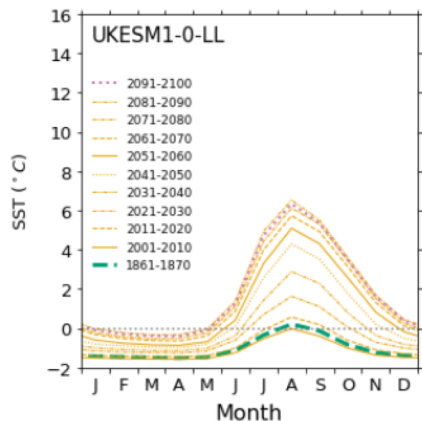
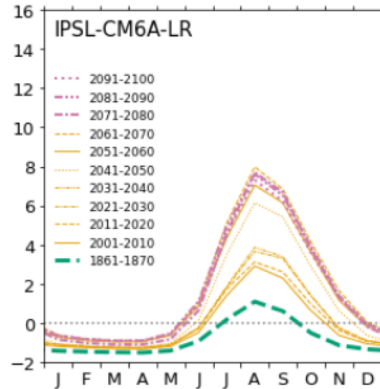
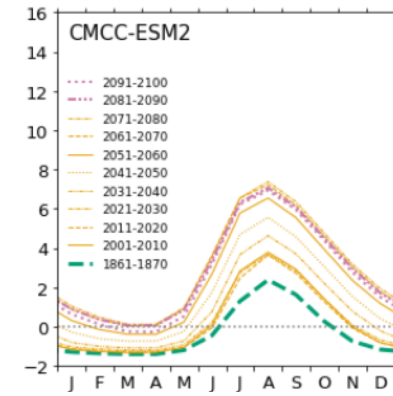
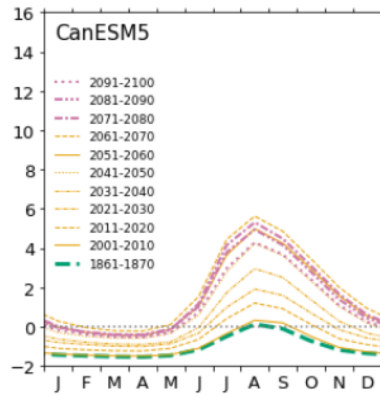
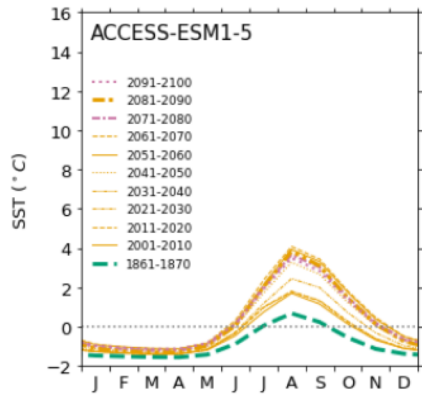


**2/5 show consistent summer  $p\text{CO}_2$  max as atmospheric  $\text{CO}_2$  declines (irreversibility)**

green - first decade of historical (1851-1860)  
 black - annual max is in winter  
 cyan - annual max in winter but also a summer max  
 yellow - summer max now annual max  
 magenta - summer max < previous decade (some reversibility)



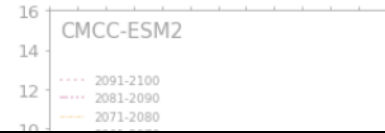
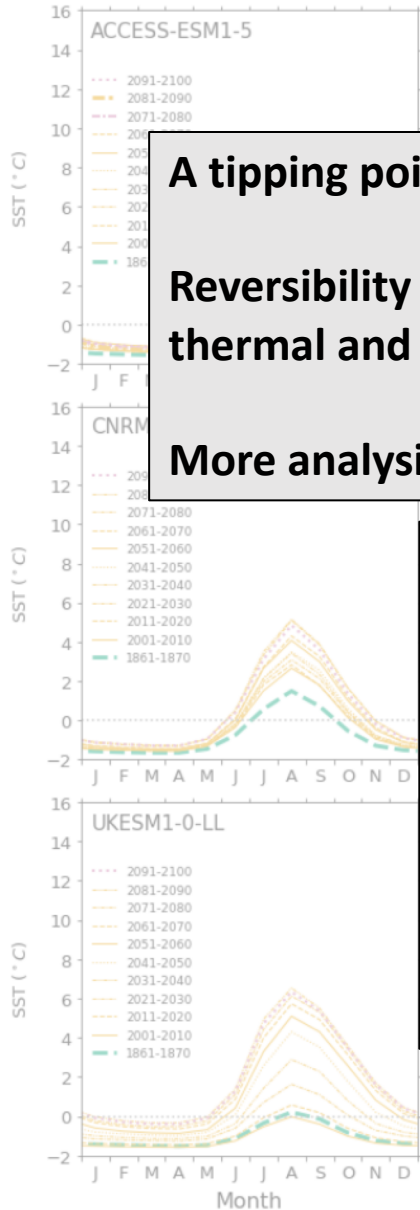
# Extent of SST amplitude declines (with declining CO<sub>2</sub>) appear to explain the reversibility



**The models with greatest SST amplitude reductions are those with greatest  $p\text{CO}_2$  reversibility**

green - first decade of historical (1851-1860)  
yellow - summer SST max > previous decade  
magenta - summer SST max < previous decade (some reversibility)

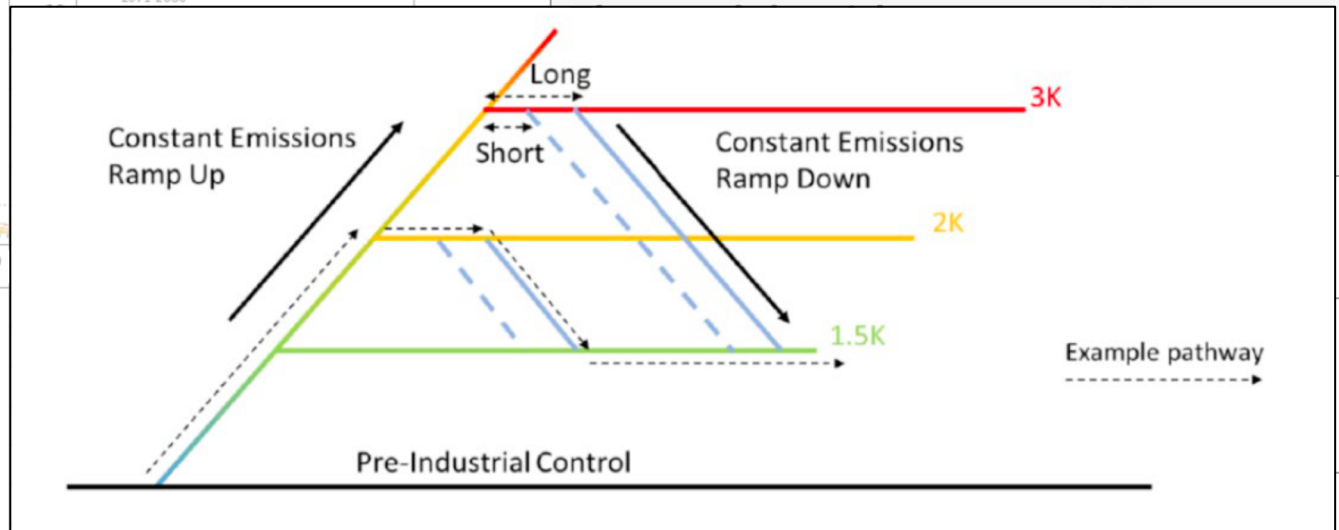
# Extent of SST amplitude declines (with declining CO<sub>2</sub>) appear to explain the reversibility



**A tipping point cascade, initiated by Arctic sea ice loss?**

**Reversibility determined by sea ice recovery and the balance between thermal and nonthermal (planktonic) drivers of the pCO<sub>2</sub> seasonal cycle?**

**More analysis required, more overshoot scenarios would be very helpful**



# Conclusions and future perspectives

**Many abrupt changes in ESM ocean biogeochemistry projections  
...but few, established tipping elements/points & limited assessment of reversibility  
and early warning indicators in ESMs**

**Arctic Ocean  $p\text{CO}_2$  seasonal cycle reorganization caused by a shift in opposing  
thermal/biological drivers under climate change**

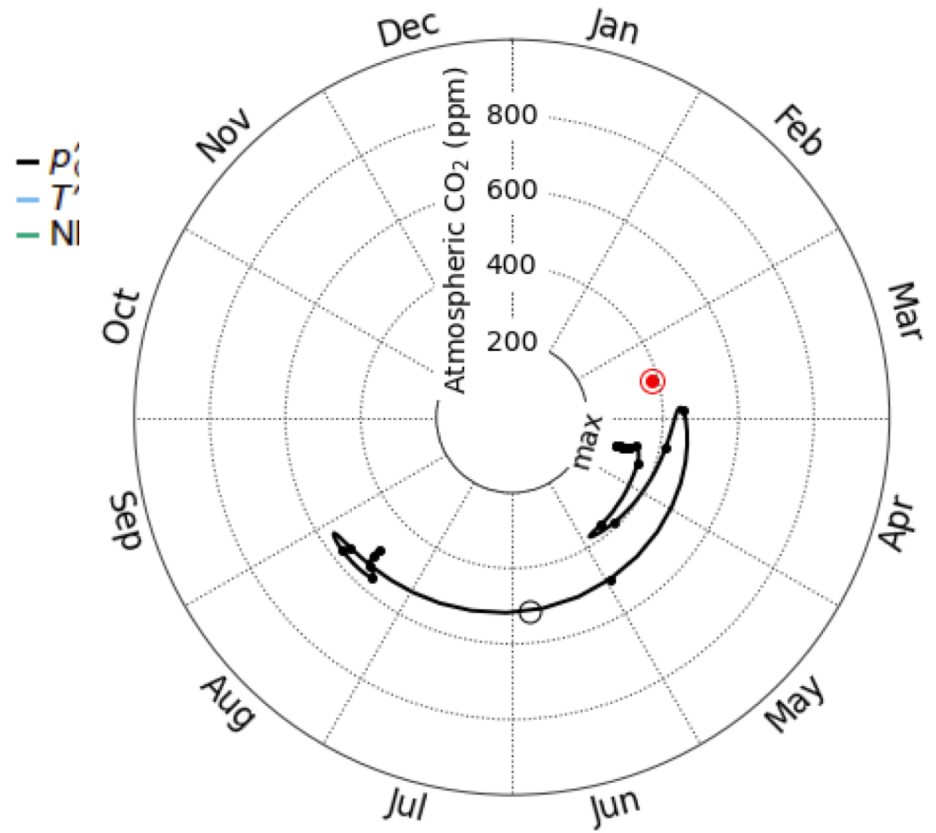
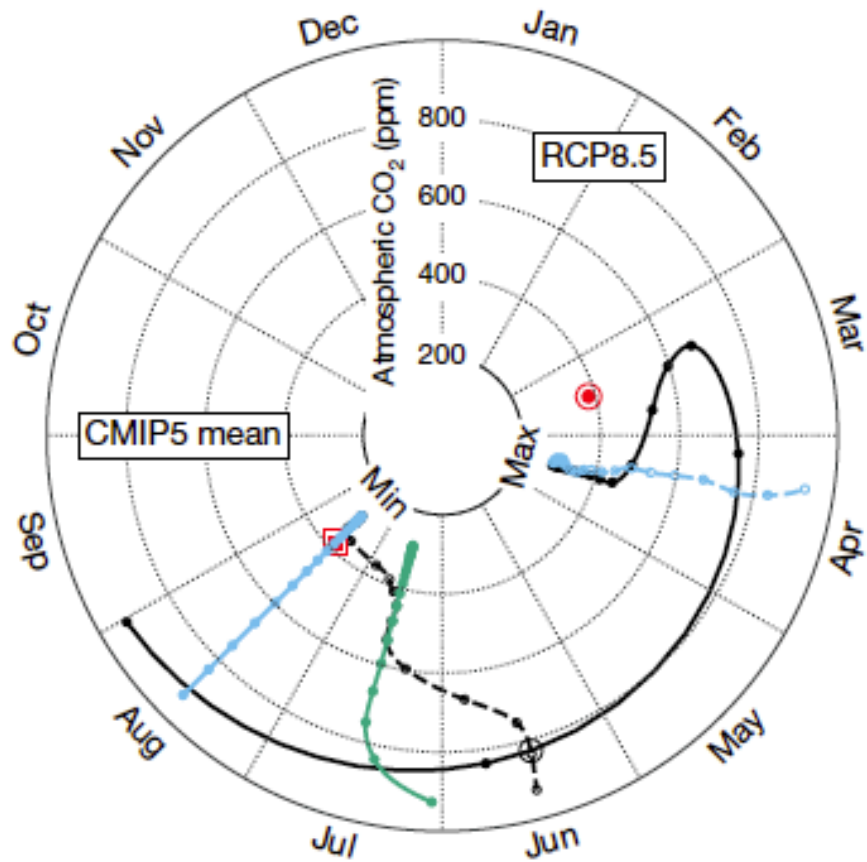
**May represent part of a sea ice driven tipping point cascade**

**Reversibility & hysteresis very model dependent in SSP5-3.4 overshoot scenario**

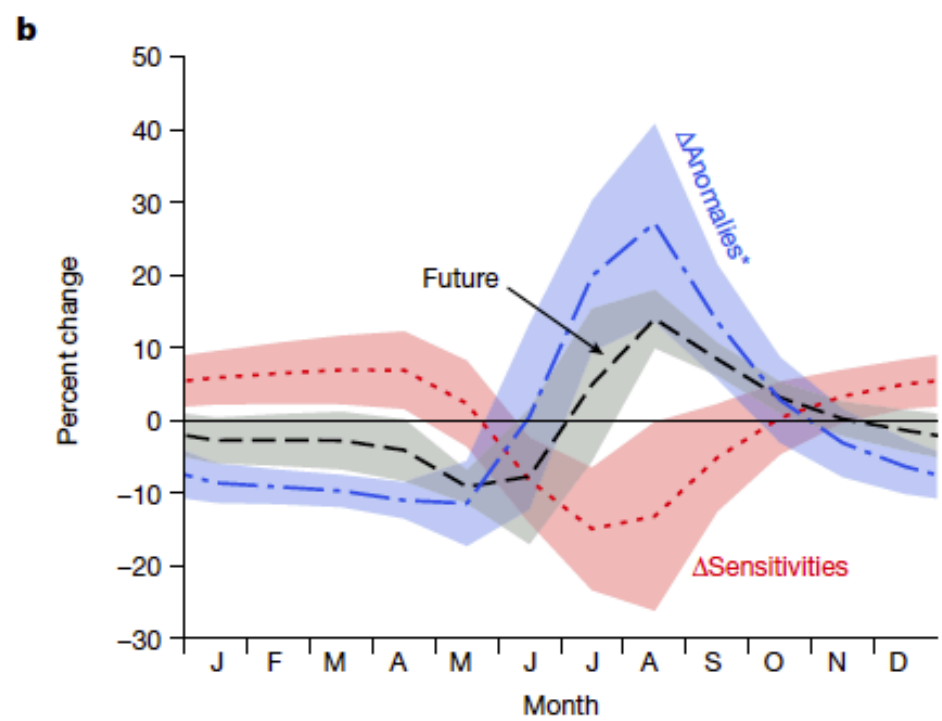
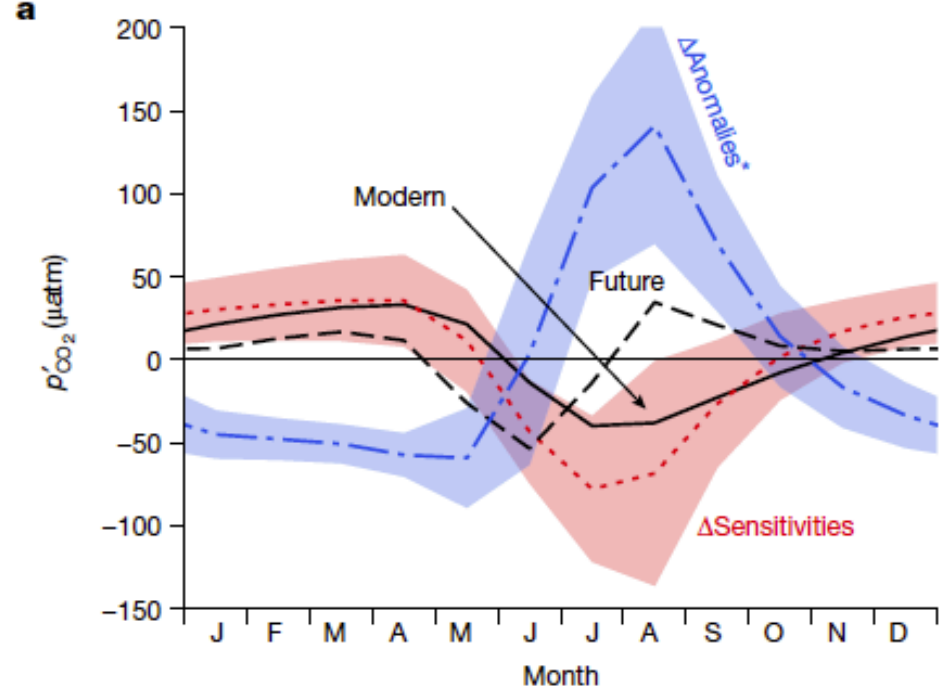
**Further OBGC tipping point assessments would benefit from harmonized protocols  
that cover a range in peak radiative forcing and negative emissions**

**New projects with tipping point focus: TIPESM, OptimESM, TIPMIP**

# Any questions?



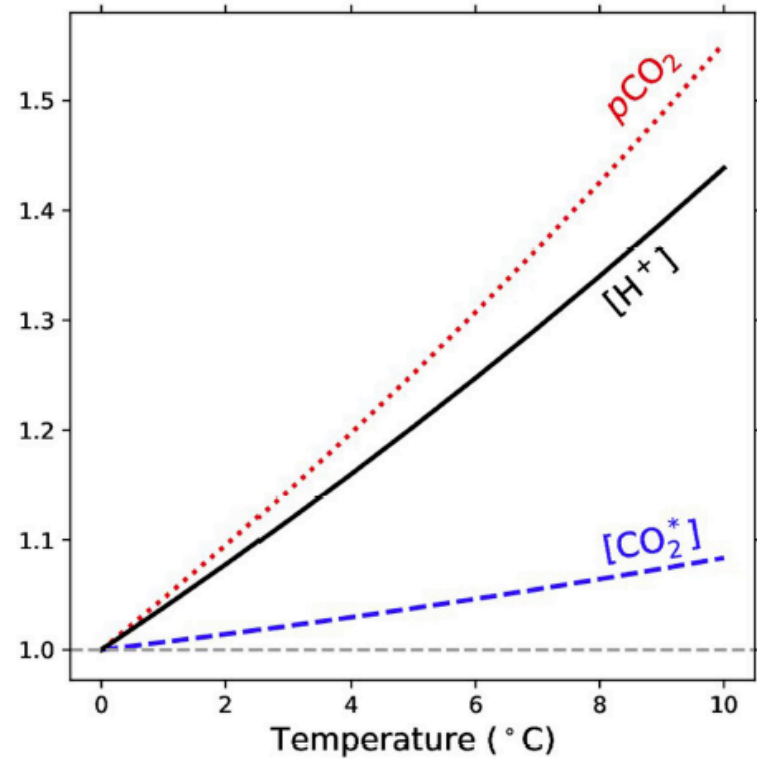
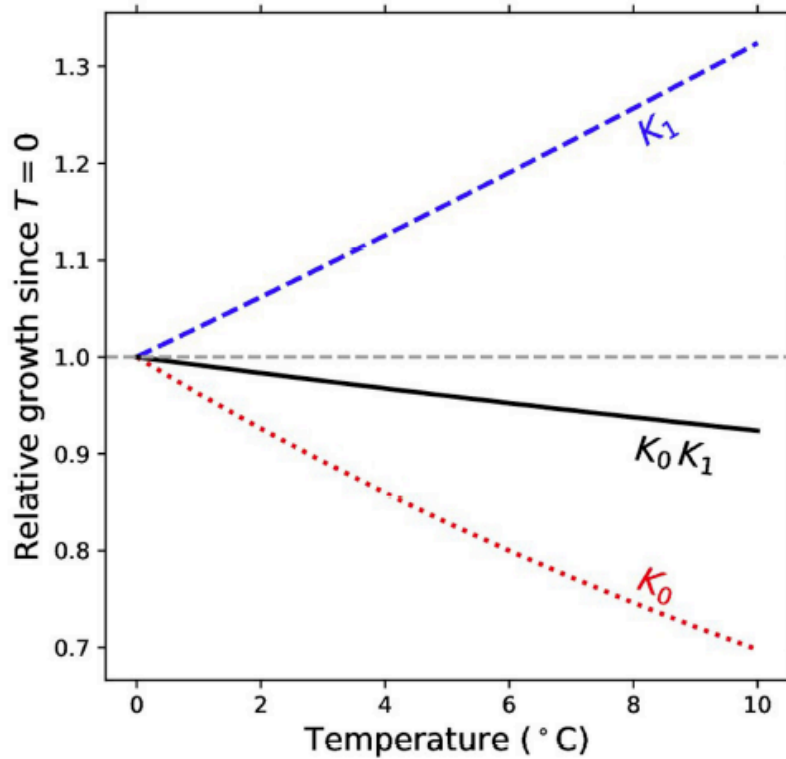
Additional slides

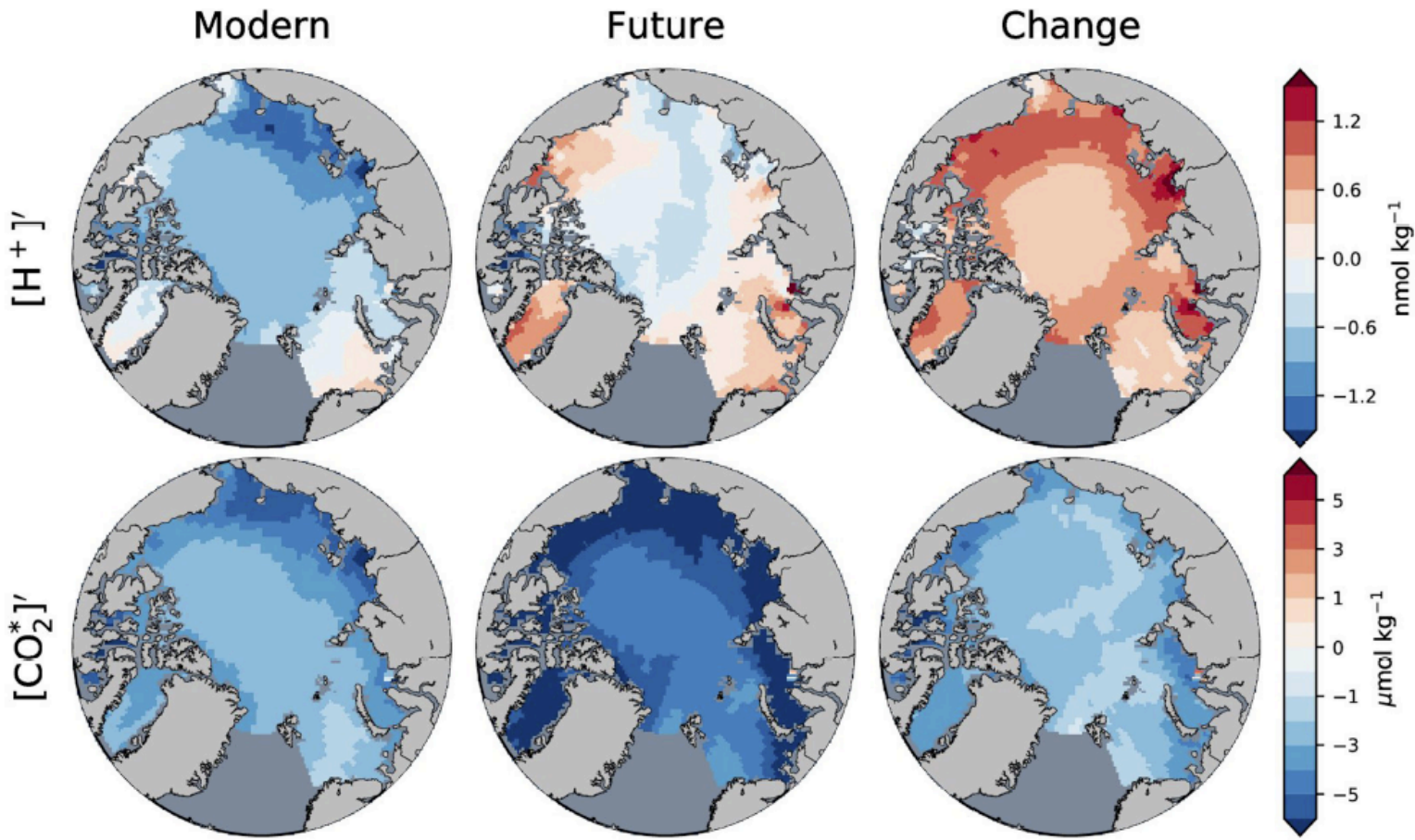


# The temperature sensitivity of $p\text{CO}_2$ , $[\text{CO}_2^*]$ and $[\text{H}^+]$

$$[\text{CO}_2^*] = K_0 C_f p_{\text{CO}_2}$$

$$p_{\text{CO}_2} = \frac{[\text{H}^+][\text{HCO}_3^-]}{C_f K_0 K_1}$$

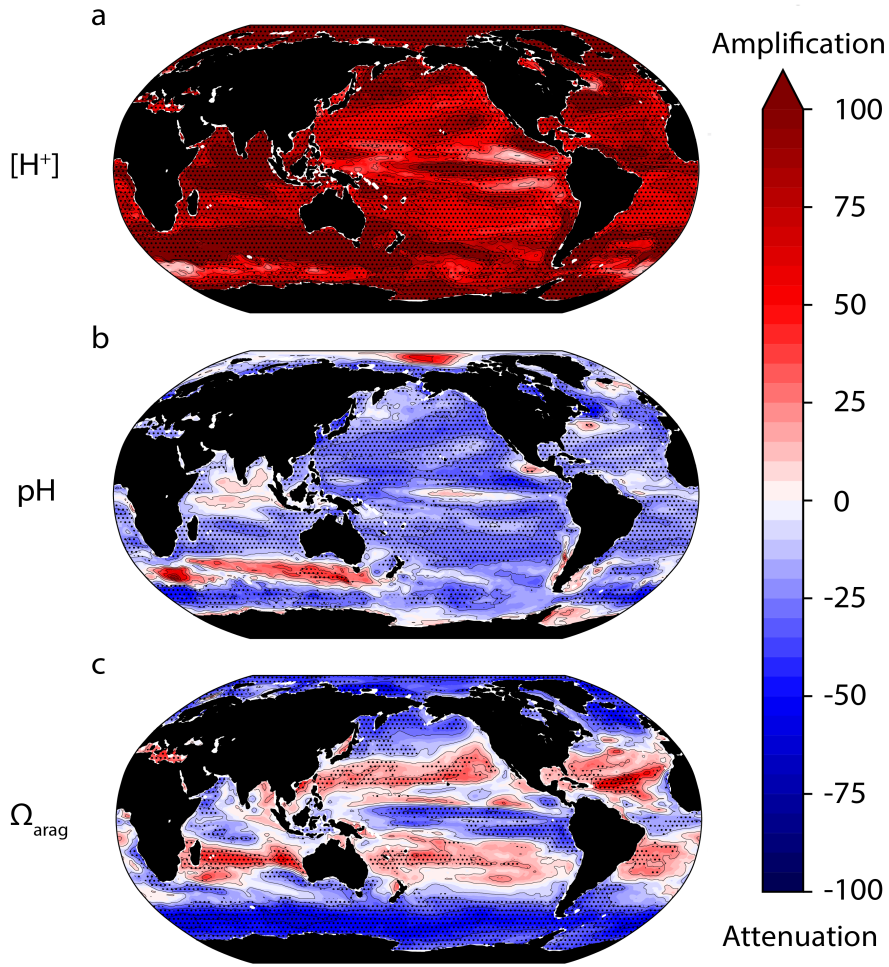






# Temporal variability: Changing seasonal amplitudes of the CO<sub>2</sub> system

## 21<sup>st</sup> century change in peak-to-peak seasonality (%)

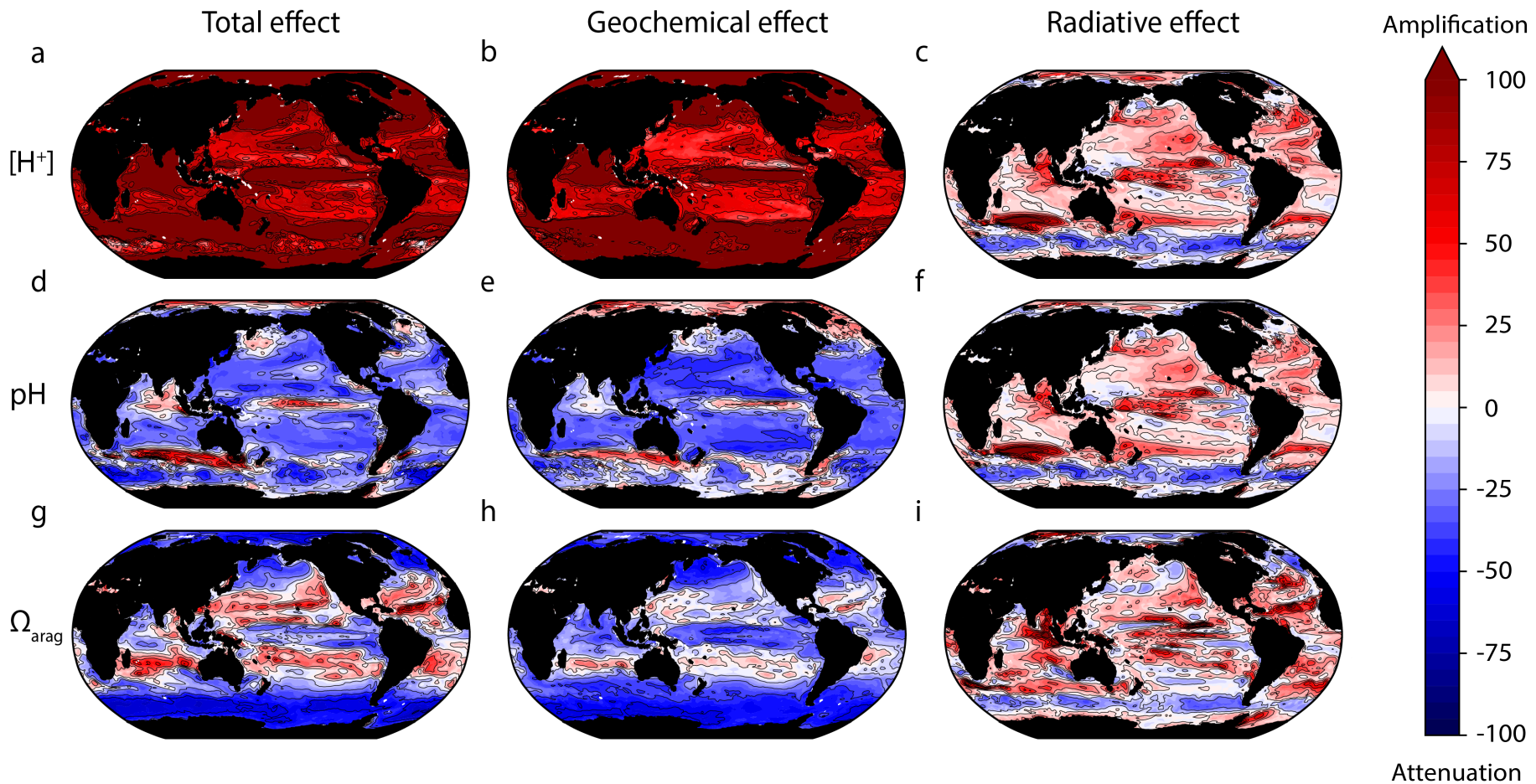


- Large **amplification** (+88%) of the [H<sup>+</sup>] seasonal cycle
  - Mild **attenuation** (-16%) of the pH seasonal cycle
  - Regional **amplification/attenuation** of the Ω<sub>arag</sub> seasonal cycle
- Supported by recent observations, Landschützer et al. (2018)

Kwiatkowski & Orr, *Nature Climate Change* (2018)  
Kwiatkowski et al., *Biogeosciences* (2020)

# Temporal variability: Changing seasonal amplitudes of the CO<sub>2</sub> system

Seasonality change driven by the geochemical effect, except for  $\Omega_{\text{arag}}$  where the radiative/climate effect dominates in the subtropics



Tipping point: *“A level of change in system properties beyond which a system reorganizes, often in a non-linear manner, and does not return to the initial state even if the drivers of the change are abated.”*

(IPCC, SROCC, 2019)