

Uncertainty quantification in numerical models, with a focus on Sensitivity Analysis

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Prévisibilité dans les sciences de l'atmosphère, des océans,
et du climat, 2-3 Octobre 2023, IHP



Outline

General introduction

Introduction to Sensitivity Analysis

Sensitivity Analysis tools

- Local Sensitivity Analysis

- Global Sensitivity Analysis

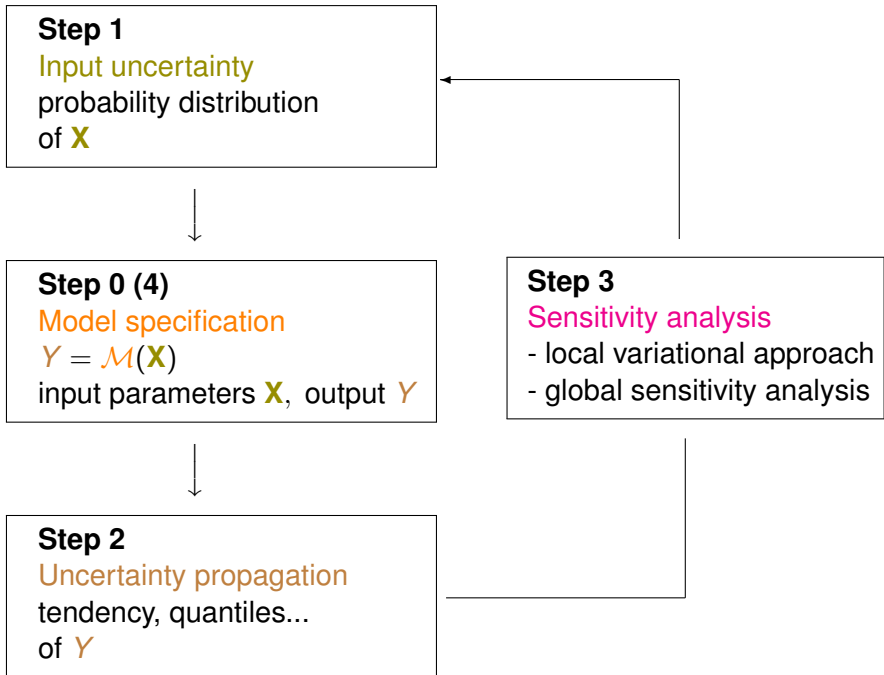
 - Screening

 - Variance-based Sensitivity Analysis

Application to MODECOGeL

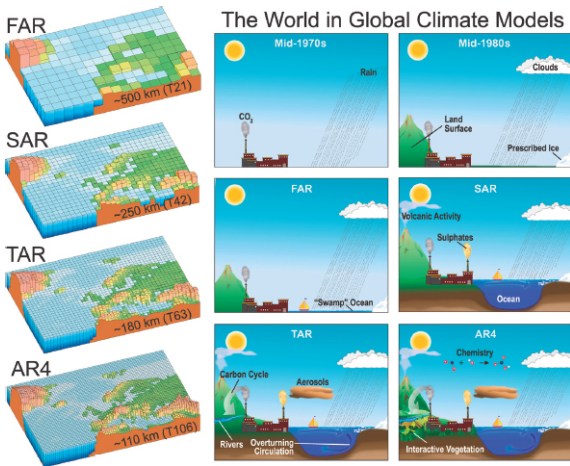
What if inputs are dependent?

Conclusion, perspectives

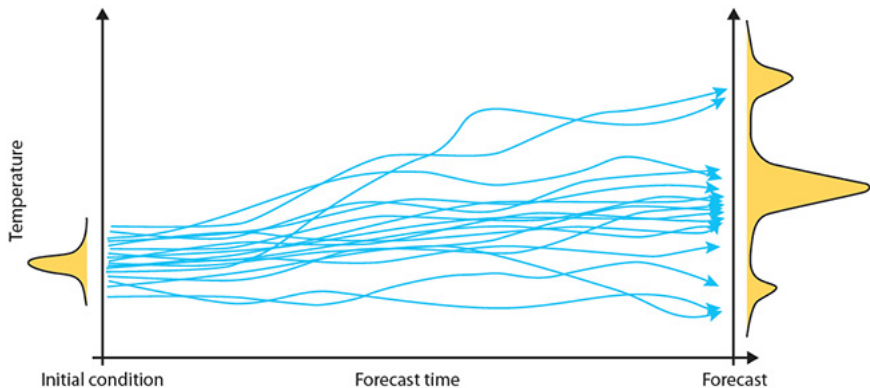


Step 0: model specification

Evolution of climate models through last century



https://www.windows2universe.org/earth/climate/climate_modeling.html



<https://www.ecmwf.int/en/about/media-centre/news/2017/twenty-five-years-ensemble-forecasting>

Step 1

Step 2

Concerning **Step 1**, uncertainty may be classified into two categories

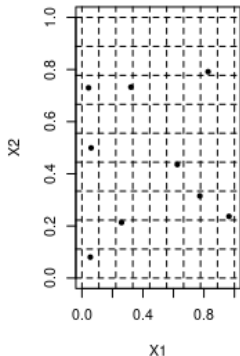
- ▶ **aleatoric** (aka **statistical**) uncertainty refers to the notion of randomness, that is, the variability in the outcome of an experiment,
- ▶ **epistemic** (aka **systematic**) uncertainty refers to uncertainty caused by a lack of knowledge.

Examples:

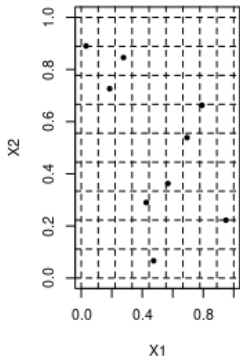
- ▶ meteorological inputs are random,
- ▶ bathymetry.

Classification is not always easy.

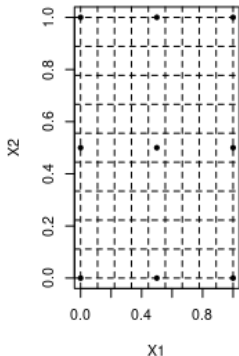
How to explore "at best" input parameter space for **Step 2** of uncertainty propagation?



Uniform design



Latin Hypercube Sampling

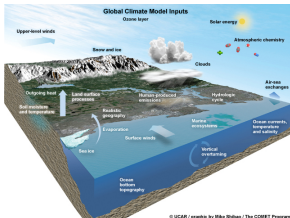


Factorial design

It is important, e.g., for parameter perturbation in view of ensemble forecast.

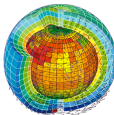
Step 1 misspecified input parameters

General introduction

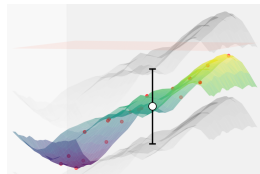


Step 0 complex models, $\mathcal{O}(10^{10})$ degrees of freedom

$$\left\{ \begin{array}{l} \frac{\partial u}{\partial t} + v \frac{\partial u}{\partial x} + w \frac{\partial u}{\partial y} + \frac{\partial u}{\partial \sigma} - f(v - v_g) + \frac{\partial \phi}{\partial x} + F_x = 0 \\ \frac{\partial v}{\partial t} + u \frac{\partial v}{\partial x} + w \frac{\partial v}{\partial y} + \frac{\partial v}{\partial \sigma} + f(u - u_g) + \frac{\partial \phi}{\partial y} + F_y = 0 \\ \frac{\partial \phi}{\partial t} + \frac{\partial \phi}{\partial \sigma} = -C_p \theta (1 + 0.85 q_v) \\ \frac{\partial p_v}{\partial t} = - \int_0^{\sigma} \left[\frac{\partial(p_v u)}{\partial x} + \frac{\partial(p_v v)}{\partial y} \right] d\sigma \\ \frac{\partial \sigma}{\partial t} = \frac{1}{p_*} \left\{ \sigma \int_0^1 \left[\frac{\partial(p_* u)}{\partial x} + \frac{\partial(p_* v)}{\partial y} \right] d\sigma - \int_0^{\sigma} \left[\frac{\partial(p_* u)}{\partial x} + \frac{\partial(p_* v)}{\partial y} \right] d\sigma \right\} \\ \frac{\partial w}{\partial t} + u \frac{\partial w}{\partial x} + v \frac{\partial w}{\partial y} + \sigma \frac{\partial w}{\partial \sigma} + \frac{\partial \sigma}{\partial t} = D \bar{w}_{rad} + F_3 + p_* = 0 \\ \frac{\partial \theta}{\partial t} + u \frac{\partial \theta}{\partial x} + v \frac{\partial \theta}{\partial y} + \sigma \frac{\partial \theta}{\partial \sigma} + F_4 + p_* = 0 \end{array} \right.$$



Step 2 uncertain predictions



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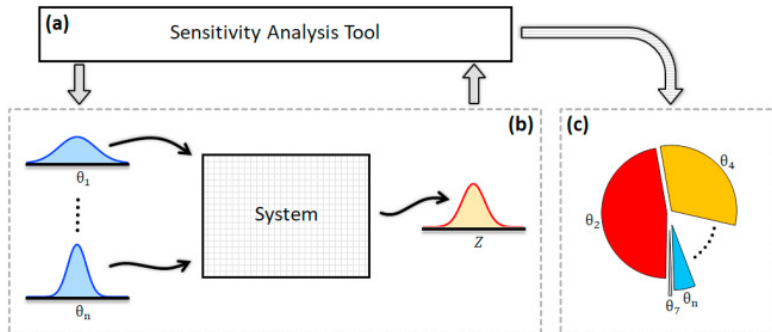
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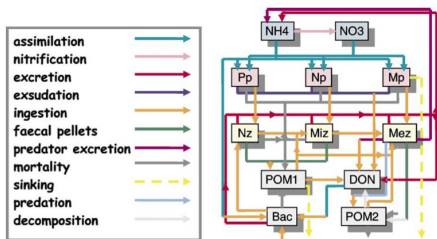
Step 3: sensitivity analysis (Razavi *et al.*, 2021)

Aim of Sensitivity Analysis: find how **model outputs** vary with **input** changes.

Application to a biogeochemical model:
ecosystem model (MODECOGeL) of the Ligurian Sea
Joint work with IGE Lab (Grenoble, FRANCE)



MODECOGeL is a one-dimensional coupled hydrodynamical-biological model.



- hydrodynamic model: 1-D vertical simplification of primitive equations for the ocean, 5 state variables;
- ecosystem model: marine biogeochemistry, 12 biological state variables.

Inputs/Outputs:

- ▷ 74 scalar input parameters;
- ▷ spatio-temporal outputs.

Main issue: calibration of the model.

Sensitivity Analysis is a preliminary step to this calibration task.

GSA for convection-permitting Numerical Weather Prediction (NWP) models Wimmer *et al* (2022)

Aim: to determine the most influential parameters on the forecast of different near-surface variables.

Model:
convective-scale AROME model.

Input parameters:

| Scheme | Parameter | Physical meaning | Default | Range |
|--------------|-----------|---|---------|------------------|
| Radiation | RSWINHF | Shortwave inhomogeneity factor | 1 | 0.6 - 1 |
| | RLWINHF | Longwave inhomogeneity factor | 1 | 0.6 - 1 |
| Microphysics | RCRIAUTI | Snow Autoconversion threshold | 0.2e-3 | 0.2e-4 - 0.25e-3 |
| | RCRIAUTC | Rain Autoconversion threshold | 1e-3 | 0.4e-3 - 1e-3 |
| | VSIGQSAT | Constant for subgrid condensation | 0.02 | 0 - 0.1 |
| Turbulence | XLINI | Minimum mixing length | 0 | 0 - 0.2 |
| | XCTD | Constant for dissipation of temperature and vapor pressure fluctuations | 1.2 | 0.98 - 1.2 |
| | XCTP | Constant for temperature and vapor pressure correlations | 4.65 | 1.035 - 22.22 |
| | XCEP | Constant for wind-pressure correlations | 2.11 | 0.225 - 4.0 |
| | XCED | Constant for dissipation of TKE | 0.85 | 0.4 - 2 |
| | XPH.LIM | Threshold value for Sc^{-1} and Pr^{-1} | 3 | 1 - 4.5 |
| | XCET | Constant for transport of TKE | 0.4 | 0.072 - 1.512 |
| Diffusion | SLHDEPSH | Strength of SLHD | 0.060 | 0.01 - 0.09 |
| | SLHDKMIN | Diffusion function minimum | 0 | -1 - 1 |
| | SLHDKMAX | Diffusion function maximum | 6 | 4 - 12 |
| Surface | XRIMAX | Critical Richardson Number | 0.2 | 0 - 0.3 |
| | XFRACZ0 | Coefficient of orographic drag | 5 | 2 - 10 |
| Convection | XCMF | Closure coefficient at bottom level | 0.065 | 0 - 0.1 |
| | XABUO | Coefficient of the buoyancy | 1 | 0.7 - 1.5 |
| | XBDETR | Coefficient of the detrainment | 1e-6 | 0 - 1 |
| | XENTR.DRY | Coefficient for dry entrainment | 0.55 | 0.1 - 0.699 |

Scalar outputs: either an averaged forecast field or a performance metric such as mean bias, root-mean square error and mean absolute error.

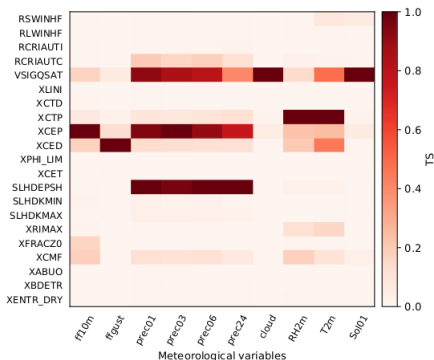
$$\text{Bias} = \frac{1}{n} \sum_{k=1}^n (y_k - o_k), \quad \text{RMSE} = \sqrt{\frac{1}{n} \sum_{k=1}^n (y_k - o_k)^2},$$

$$\text{MAE} = \frac{1}{n} \sum_{k=1}^n |y_k - o_k|.$$

These scores are computed using SYNOP (surface synoptic observations) and the French real-time meteorological observation network (Tardieu & Leroy, 2003).

Notation: n is the number of in-situ measures, y_k is the model output and o_k is the k -th observation.

On the figure below (Wimmer *et al*, 2022), outputs are spatial-averaged scores computed for 10-meter wind speed **ff10m**, 10-meter wind gust **ffgust**, 1-hourly, 3-hourly, 6-hourly and 24-hourly accumulated precipitation **prec01**, **prec03**, **prec06**, **prec24**, total cloud cover **cloud**, 2-meter relative humidity **RH2m**, 2-meter temperature **T2m** and 1-hourly downward global solar radiation **Sol01**. Total Sobol' indices measure the sensitivity of each of these outputs with respect to **input parameters**.



The higher the Total Sobol' Index is, the darker the colour is, the more influential the parameter is.

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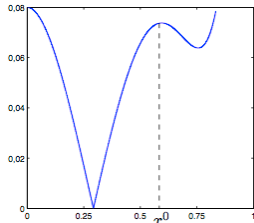
$$\mathcal{M} : \begin{cases} \mathbb{R}^d & \rightarrow \mathbb{R} \\ \mathbf{x} & \mapsto y = \mathcal{M}(x_1, \dots, x_d) \end{cases}$$

Local Sensitivity Analysis is based on Taylor approximation:

$$\mathcal{M}(\mathbf{x}) \approx \mathcal{M}(\mathbf{x}^0) + \sum_{i=1}^d \left(\frac{\partial \mathcal{M}}{\partial x_i} \right)_{\mathbf{x}^0} (x_i - x_i^0).$$

First order sensitivity index for

input i : $\left(\frac{\partial \mathcal{M}}{\partial x_i} \right)_{\mathbf{x}^0}$.



Pros: Low computational cost even for large d if one uses the adjoint stae method (see, e.g., Plessix, 2006).

Cons: Local analysis, not well-suited for highly nonlinear models.

Global Sensitivity Analysis with Screening

Main objective: to screen among a large amount of inputs which ones are non influential on the quantity of interest (QoI).

Advantages: moderate computational cost.

Drawbacks: partial information, no hierarchisation.

A OAT screening method : Morris, 1991

OAT One At a Time we vary the factors one by one.

The screening method proposed by Morris is a global OAT approach.

Model $Y = \mathcal{M}(\mathbf{X})$, $\mathbf{X} = (X_1, \dots, X_d)$ with the X_i s independent uniform random variables on $[0, 1]$.

More details on the method :

- input discretization on a grid with p values $\{0, \frac{1}{p-1}, \dots, 1\}$.
- Δ a multiple of $1/(p-1)$, fixed once for all.
- $\Omega := \{0, \frac{1}{p-1}, \dots, 1\}^d$.
- $\Omega_i^\Delta := \{\mathbf{x} \in \Omega \text{ such that } (x_1, \dots, x_{i-1}, x_i + \Delta, x_{i+1}, \dots, x_d) \in \Omega\}$.

Definition

Elementary effect of X_i computed at $\mathbf{x} \in \Omega_i^\Delta$,

$$d_i(\mathbf{x}) = \frac{1}{\Delta} \{ \mathcal{M}(x_1, \dots, x_{i-1}, x_i + \Delta, x_{i+1}, \dots, x_d) - \mathcal{M}(\mathbf{x}) \} .$$

There are $p^{d-1}(p - \Delta(p-1))$ elementary effects to compute.

Steps :

- ▶ one draws uniformly a r -sample in $\Omega_i^\Delta : \mathbf{x}^1, \dots, \mathbf{x}^r$;
- ▶ one computes $d_i(\mathbf{x}^j)$, $j = 1, \dots, r$, $i = 1, \dots, d$;
- ▶ one computes

$$\mu_i = \frac{1}{r} \sum_{j=1}^r d_i(\mathbf{x}^j), \quad \sigma_i^2 = \frac{1}{r} \sum_{j=1}^r (d_i(\mathbf{x}^j) - \mu_i)^2.$$

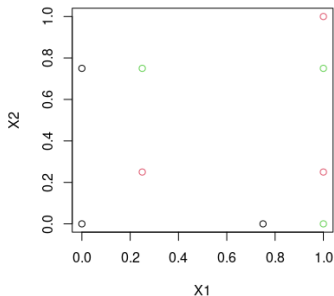
| | σ_i^2 low | σ_i^2 high |
|----------------|------------------|------------------------------------|
| $ \mu_i $ low | non influential | nonlinearities and/or interactions |
| $ \mu_i $ high | influential | nonlinearities and/or interactions |

Elementary effect of X_i computed at $\mathbf{x} \in \Omega_i^\Delta$,

$$d_i(\mathbf{x}) = \frac{1}{\Delta} \{ \mathcal{M}(x_1, \dots, x_{i-1}, x_i + \Delta, x_{i+1}, \dots, x_d) - \mathcal{M}(\mathbf{x}) \} .$$

The **efficiency** of the method "number of elementary effects computed / number of model runs" is equal to $1/2$.

Morris (1991) presents an adaptation with an efficiency equal to $d/(d+1)$, with d the input space dimension.



$r = 3$ Morris trajectories with $p = 5$,
 $\Delta = 3/4$.

A toy example Advection-reaction-diffusion equation with Dirichlet boundary condition :

$$\left\{ \begin{array}{l} \frac{\partial u}{\partial t} = -r \cdot u - a \frac{\partial u}{\partial x} + \lambda \frac{\partial^2 u}{\partial x^2} + f \quad x \in [0, L], t \in [0, T] \\ u(x=0, t) = \psi_1(t) \quad t \in [0, T] \\ u(x=L, t) = \psi_2(t) \quad t \in [0, T] \\ u(x, t=0) = g(x) \quad x \in (0, L). \end{array} \right.$$

A : energy norm of the solution at time $t = T$.

Sensitivity of **A** with respect to (a, r, λ) ? Uncertain input parameters are modeled as $a, r \sim \mathcal{U}([0.4, 0.6])$, $\lambda \sim \mathcal{U}([0.04, 0.06])$.

Scheme : 2-steps Adams-Moulton, sample size equals 2^{13} .

Sensitivity measures based on variance : $S_a = 0.0188$, $S_\lambda = 0.7299$, $S_r = 0.2488$, $S_a + S_\lambda + S_r = 0.988$.

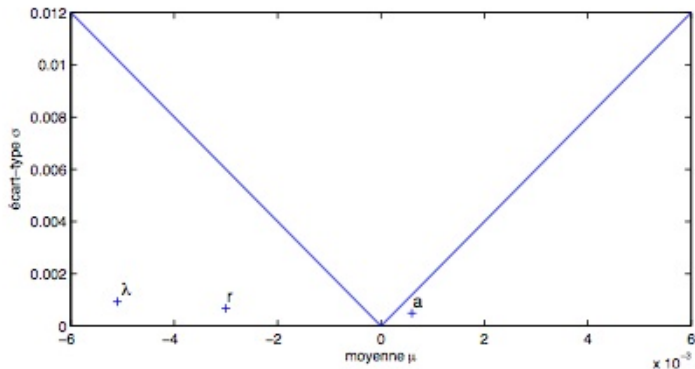


Figure: Morris with $p = 50$, $\Delta = 25/49$.

Variance-based Global Sensitivity Analysis

Independent framework: $P(d\mathbf{x}) = P_1(dx_1) \dots P_d(dx_d)$

Let

$$\mathcal{M}: \begin{cases} \mathbb{R}^d & \rightarrow \mathbb{R} \\ \mathbf{x} = (x_1, \dots, x_d) & \mapsto y = \mathcal{M}(\mathbf{x}) \end{cases}$$

Does the output Y vary more or less when fixing one of its **input parameters**? $V[Y|X_i = x_i]$, how to choose x_i ?

$$\rightarrow E[V(Y|X_i)] = V[Y] - V[E(Y|X_i)].$$

First-order Sobol' indices: $0 \leq S_i = \frac{V[E(Y|X_i)]}{V[Y]} \leq 1.$

The more this quantity is close to 1, the more fixing X_i reduces the variance of Y : the input X_i is influential.

More generally,

$$S_i = \frac{V[E[Y|X_i]]}{V[Y]}, \quad 1 \leq i \leq d$$

$$S_{i,j} = \frac{V[E[Y|X_i, X_j]] - V[E[Y|X_i]] - V[E[Y|X_j]]}{V[Y]}, \quad 1 \leq i \neq j \leq d \dots$$

We have $1 = \sum_{i=1}^d S_i + \sum_{i \neq j} S_{i,j} + \dots + S_{1,\dots,d}$

Factors Prioritization (FP): which factor should one try to determine first to get the largest expected reduction in the variance of the model output? \rightarrow first order Sobol' indices do the job.

Total Sobol' indices:

$$i = 1, \dots, d \quad S_i^{\text{tot}} = \sum_{\mathbf{u} \subseteq \{1, \dots, d\}, \mathbf{u} \cap \{i\} \neq \emptyset} S_{\mathbf{u}}$$

Factors Fixing (FF): which input factors can be fixed, anywhere in their range of variation, without sensibly affecting a specific output of interest? \rightarrow total Sobol' indices do the job.

We have:

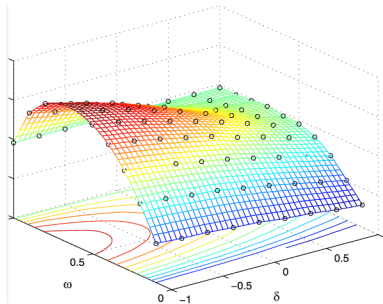
$$S_i^{\text{tot}} = \frac{E[V[Y|\mathbf{X}_{-i}]]}{V[Y]} = 1 - \frac{V[E[Y|\mathbf{X}_{-i}]]}{V[Y]}$$

with $\mathbf{X}_{-i} = (X_1, \dots, X_{i-1}, X_{i+1}, \dots, X_d)$.

Sobol' index estimation

- ▶ Sobol' indices can be estimated from input/output samples $(\mathbf{x}^{(i)}, Y^{(i)} = \mathcal{M}(\mathbf{x}^{(i)}))$, $1 \leq i \leq n$.

- ▶ Metamodels can be built to speed up computations.



Some additional issues

- ▶ visualization for complex outputs,
- ▶ correlated inputs,
- ▶ stochastic models. . .

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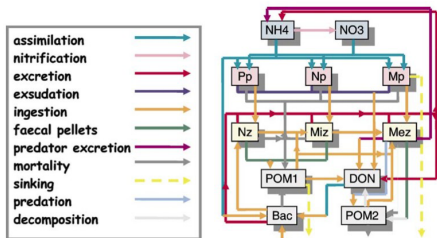
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see Prieur *et al*, 2019



MODECOGeL is a one-dimensional coupled hydrodynamical-biological model.



- hydrodynamic model: 1-D vertical simplification of primitive equations for the ocean, 5 state variables;
- ecosystem model: marine biogeochemistry, 12 biological state variables.

▷ 74 independent scalar parameters

| Index | Name | Unit | Pdf | Mean | Std | Std/Mean |
|-------|---|----------|------------------------|-------|--------|----------|
| 1 | PicP max growth rate | t^{-1} | $\Gamma(25, 0.12)$ | 3. | 0.6 | 20% |
| 2 | NanP max growth rate | t^{-1} | $\Gamma(25, 0.1)$ | 2.5 | 0.5 | 20% |
| 3 | MicP max growth rate | t^{-1} | $\Gamma(25, 0.08)$ | 2. | 0.4 | 20% |
| 4 | dependence of NO ₃ limitation to NH ₄ | C^{-1} | $\Gamma(0.0, 0.00025)$ | 1.46 | 0.0013 | 1% |
| 5 | NO ₃ semi-saturation for PicP | C | $\Gamma(4, 0.125)$ | 0.5 | 0.25 | 50% |
| 6 | NO ₃ semi-saturation for NanP | C | $\Gamma(4, 0.175)$ | 0.7 | 0.35 | 50% |
| 7 | NO ₃ semi-saturation for MicP | C | $\Gamma(4, 0.25)$ | 1.0 | 0.5 | 50% |
| 8 | NH ₄ semi-saturation for PicP | C | $\Gamma(4, 0.075)$ | 0.3 | 0.15 | 50% |
| 9 | NH ₄ semi-saturation for NanP | C | $\Gamma(4, 0.125)$ | 0.5 | 0.25 | 50% |
| 10 | NH ₄ semi-saturation for MicP | C | $\Gamma(4, 0.175)$ | 0.7 | 0.35 | 50% |
| 11 | optimal PAR for PicP | J | $\Gamma(25, 0.4)$ | 10. | 2. | 20% |
| 12 | optimal PAR for NanP | J | $\Gamma(25, 0.6)$ | 15. | 3. | 20% |
| 13 | optimal PAR for MicP | J | $\Gamma(25, 0.8)$ | 20. | 4. | 20% |
| 14 | variation of light limitation for PicP | — | $-\Gamma(4, 0.2)$ | -0.8 | 0.4 | 50% |
| 15 | variation of light limitation for NanP | — | $-\Gamma(4, 0.175)$ | -0.7 | 0.35 | 50% |
| 16 | variation of light limitation for MicP | — | $-\Gamma(4, 0.15)$ | -0.6 | 0.3 | 50% |
| 17 | optimal temperature for PicP | T | $N(15, 3^2)$ | 15. | 3. | 20% |
| 18 | optimal temperature for NanP | T | $N(15, 3^2)$ | 15. | 3. | 20% |
| 19 | optimal temperature for MicP | T | $N(16, 3.2^2)$ | 16. | 3.2 | 20% |
| 20 | variation of temp. limitation for PicP | — | $-\Gamma(4, 0.125)$ | -0.5 | 0.25 | 50% |
| 21 | variation of temp. limitation for NanP | — | $-\Gamma(4, 0.125)$ | -0.5 | 0.25 | 50% |
| 22 | variation of temp. limitation for MicP | — | $-\Gamma(4.0, 0.1375)$ | -0.55 | 0.275 | 50% |
| 23 | bacteria growth limitation | — | $\Gamma(4, 0.15)$ | 0.6 | 0.3 | 50% |
| 24 | semi-saturation for BAC growth | C | $\Gamma(4, 0.125)$ | 0.5 | 0.25 | 50% |
| 25 | exclusion ratio for PicP | — | $\Gamma(4, 0.015)$ | 0.06 | 0.03 | 50% |
| 26 | exclusion ratio for NanP | — | $\Gamma(4, 0.0125)$ | 0.05 | 0.025 | 50% |
| 27 | exclusion ratio for MicP | — | $\Gamma(4, 0.01)$ | 0.04 | 0.02 | 50% |
| 28 | max ingestion rate for NanZ | t^{-1} | $\Gamma(25, 0.12)$ | 3. | 0.6 | 20% |
| 29 | max ingestion rate for MicZ | t^{-1} | $\Gamma(25, 0.08)$ | 2. | 0.4 | 20% |
| 30 | max ingestion rate for MesZ | t^{-1} | $\Gamma(25, 0.06)$ | 1.5 | 0.3 | 20% |
| 31 | threshold ingestion for NanZ | C | $\Gamma(4, 0.0125)$ | 0.05 | 0.025 | 50% |
| 32 | threshold ingestion for MicZ | C | $\Gamma(4, 0.0075)$ | 0.03 | 0.015 | 50% |
| 33 | threshold ingestion for MesZ | C | $\Gamma(4, 0.0025)$ | 0.01 | 0.005 | 50% |
| 34 | semi-saturation for ingestion by NanZ | C | $\Gamma(4, 0.125)$ | 0.5 | 0.25 | 50% |
| 35 | semi-saturation for ingestion by MicZ | C | $\Gamma(4, 0.1875)$ | 0.75 | 0.375 | 50% |
| 36 | semi-saturation for ingestion by MesZ | C | $\Gamma(4, 0.25)$ | 1. | 0.5 | 50% |
| 37 | efficiency of MesZ on BAC | — | $\beta(4.2, 1.05)$ | 0.8 | 0.16 | 20% |
| 38 | efficiency of NanZ on BAC | — | $\beta(4.2, 1.05)$ | 0.8 | 0.16 | 20% |
| 39 | efficiency of MicZ on NanZ | — | $\beta(4.2, 1.05)$ | 0.8 | 0.16 | 20% |
| 40 | efficiency of MesZ on MicZ | — | $\beta(4.2, 1.05)$ | 0.8 | 0.16 | 20% |
| 41 | efficiency of MicZ on MOP1 | — | $\beta(19.8, 79.2)$ | 0.2 | 0.04 | 20% |
| 42 | efficiency of MesZ on MOP1 | — | $\beta(19.8, 79.2)$ | 0.2 | 0.04 | 20% |
| 43 | efficiency of MesZ on MOP2 | — | $\beta(19.8, 79.2)$ | 0.2 | 0.04 | 20% |
| 44 | mortality rate for PicP | t^{-1} | $\Gamma(4, 0.015)$ | 0.06 | 0.03 | 50% |
| 45 | mortality rate for NanP | t^{-1} | $\Gamma(4, 0.0125)$ | 0.05 | 0.025 | 50% |
| 46 | mortality rate for MicP | t^{-1} | $\Gamma(4, 0.01)$ | 0.04 | 0.02 | 50% |
| 47 | mortality rate for NanZ | t^{-1} | $\Gamma(4, 0.015)$ | 0.06 | 0.03 | 50% |
| 48 | mortality rate for MicZ | t^{-1} | $\Gamma(4, 0.0125)$ | 0.05 | 0.025 | 50% |
| 49 | mortality rate for MesZ | t^{-1} | $\Gamma(4, 0.0075)$ | 0.03 | 0.015 | 50% |
| 50 | mortality rate for BAC | t^{-1} | $\Gamma(4, 0.015)$ | 0.06 | 0.03 | 50% |
| 51 | threshold for predation | C | $\Gamma(4, 0.005)$ | 0.02 | 0.01 | 50% |
| 52 | maximum predation rate on MesZ | Γ | $\Gamma(4, 0.25)$ | 1. | 0.5 | 50% |
| 53 | semi-saturation for predation on MesZ | C | $\Gamma(4, 0.25)$ | 1. | 0.5 | 50% |
| 54 | excreted fraction of predation on MesZ | — | $\beta(2.33, 4.67)$ | 0.333 | 0.167 | 50% |
| 55 | fraction of grazing used for growth of NanZ | — | $\beta(4.2, 1.05)$ | 0.8 | 0.16 | 20% |
| 56 | fraction of grazing used for growth of MicZ | — | $\beta(4.2, 1.05)$ | 0.8 | 0.16 | 20% |
| 57 | fraction of grazing used for growth of MesZ | — | $\beta(4.2, 1.05)$ | 0.8 | 0.16 | 20% |
| 58 | fraction of POM used for growth of MicZ | — | $\beta(12, 12)$ | 0.5 | 0.1 | 20% |
| 59 | fraction of POM used for growth of MesZ | — | $\beta(12, 12)$ | 0.5 | 0.1 | 20% |
| 60 | excretion rate for NanZ | t^{-1} | $\Gamma(4, 0.0375)$ | 0.15 | 0.075 | 50% |
| 61 | excretion rate for MicZ | t^{-1} | $\Gamma(4, 0.025)$ | 0.1 | 0.05 | 50% |
| 62 | excretion rate for MesZ | t^{-1} | $\Gamma(4, 0.0125)$ | 0.05 | 0.025 | 50% |
| 63 | excretion rate for BAC | t^{-1} | $\Gamma(4, 0.0375)$ | 0.15 | 0.075 | 50% |
| 64 | temperature variation of excretion for NanZ | — | LogGamma | 1.05 | 0.0525 | 5% |
| 65 | temperature variation of excretion for MicZ | — | LogGamma | 1.05 | 0.0525 | 5% |
| 66 | temperature variation of excretion for MesZ | — | LogGamma | 1.02 | 0.061 | 5% |
| 67 | temperature variation of excretion for BAC | — | LogGamma | 1.04 | 0.052 | 5% |
| 68 | fraction of excretion as DOM | — | $\beta(2.75, 8.25)$ | 0.25 | 0.125 | 50% |
| 69 | POM1 decomposition rate | t^{-1} | $\Gamma(4, 0.0125)$ | 0.065 | 0.0325 | 50% |
| 70 | POM2 decomposition rate | t^{-1} | $\Gamma(4, 0.015)$ | 0.06 | 0.03 | 50% |
| 71 | sedimentation velocity for MicP | V | $\Gamma(4, 0.25)$ | 1. | 0.5 | 50% |
| 72 | nitrification rate | t^{-1} | $\Gamma(4, 0.0075)$ | 0.03 | 0.015 | 50% |
| 73 | light attenuation coefficient in sea water | — | $\Gamma(25, 0.0016)$ | 0.04 | 0.008 | 20% |
| 74 | fraction of photosynthetically active radiation | — | $\Gamma(25, 0.02)$ | 0.5 | 0.1 | 20% |

State variables

The ecosystem model provides a 12-component description of the ecosystem of the Ligurian Sea.

| Variable | Acronym | Name |
|----------|-------------|-------------------------------------|
| C_1 | NO3 | Nitrate |
| C_2 | NH4 | Ammonium |
| C_3 | PicP | Picophytoplankton |
| C_4 | NanP | Nanophytoplankton |
| C_5 | MicP | Microphytoplankton |
| C_6 | NanZ | Nanozooplankton |
| C_7 | MicZ | Microzooplankton |
| C_8 | MesZ | Mesozooplankton |
| C_9 | BAC | Bacteria |
| C_{10} | DON | Dissolved organic nitrogen |
| C_{11} | POM1 | Particulate organic matter (size 1) |
| C_{12} | POM2 | Particulate organic matter (size 2) |

The time evolution of each state variable is governed by the equation:

$$\frac{\partial C_i}{\partial t} = \text{ADV}_i + \text{DIFF}_i + \text{SMS}_i \quad \text{with} \quad \text{SMS}_i = \sum_{j \neq i} \text{FLUX}(C_j \rightarrow C_i)$$

where ADV_i and DIFF_i are advection and diffusion terms, and SMS_i is the “source minus sink” term summing up the fluxes ($\text{FLUX}(C_j \rightarrow C_i)$) between the various components of the ecosystem. We also introduce chlorophyll concentration $C_0 = \alpha(C_3 + C_4 + C_5)$.

Qols

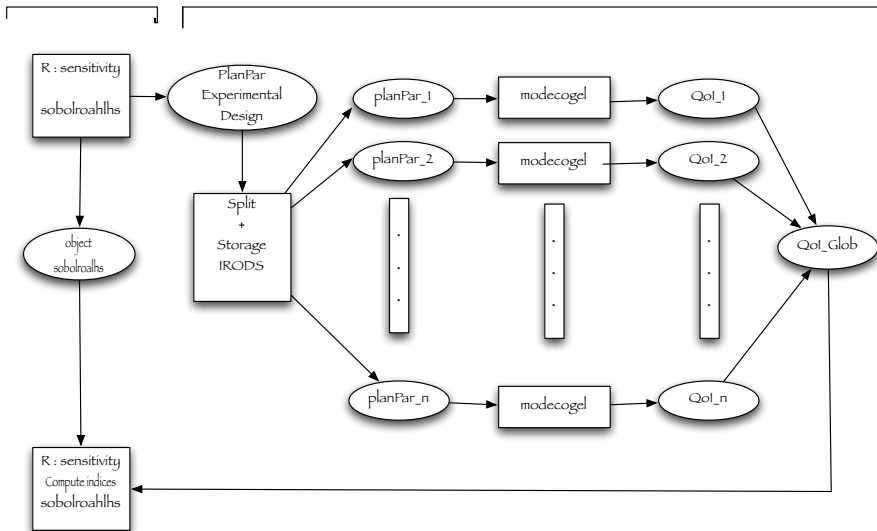
| Index j | Name | Definition |
|-----------|-------------------------------------|---|
| 1 | surface maximum | $\max_t C_i(0, t)$ |
| 2 | time of surface maximum | $\operatorname{argmax}_t C_i(0, t)$ |
| 3 | maximum of vertical average | $\max_t \frac{1}{Z} \int_0^Z C_i(z, t) dz$ |
| 4 | time of maximum of vertical average | $\operatorname{argmax}_t \frac{1}{Z} \int_0^Z C_i(z, t) dz$ |
| 5 | time and vertical average | $\frac{1}{ZT} \int_0^T \int_0^Z C_i(z, t) dz dt$ |

Quantities of interest Y_{ij} . The maximum depth for averaging is $Z = 40$ m, and T is the total duration of the experiment.

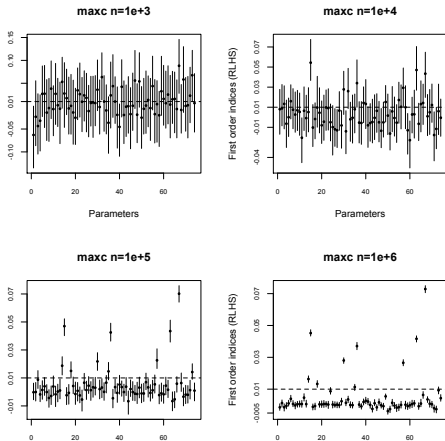
Processing chain

Workstation/laptop

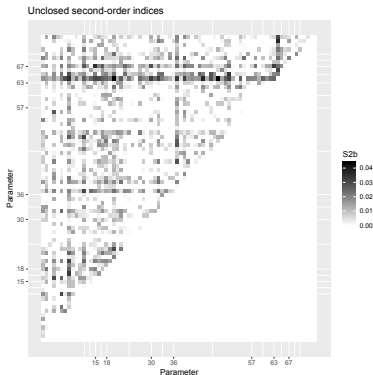
Local Grid Computing / Distributed Storage



How the results look like?



Estimated first-order indices (y -axis) with their 95% confidence interval for the 74 model parameters (x -axis), for $n = 10^3, 10^4, 10^5$ and 10^6 , in the case of the output Y_{01} . The dashed horizontal line corresponds to a threshold arbitrarily chosen to be 0.01. Confidence intervals were obtained with a bootstrap procedure and a bootstrap sample size of 100.



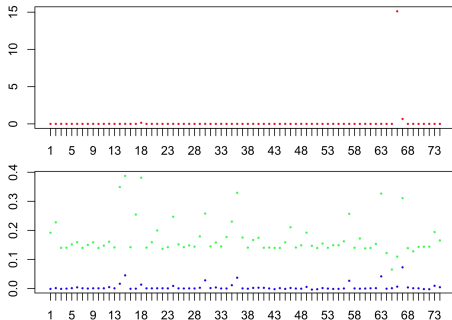
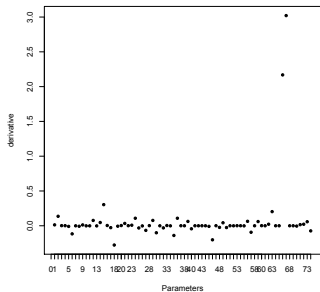
Map (74×74) of the second-order unclosed Sobol indices for QoI Y_{01} . The x and y axes correspond to the number of the parameters, and the grey scale to the value of the index. Note that the numbers indicated on the axes correspond to parameters with high first-order indices.

Top eight ranking of the local derivative $\partial Y / \partial X_j$, and first-order and total Sobol' indices S_j and S_j^{tot} .

| j | 2 | 14 | 15 | 18 | 30 | 35 | 36 | 46 | 57 | 63 | 66 | 67 |
|-----------------------------|-----------------|-----------------|-----------------|-----------------|-----------------|-----------------|-----------------|-----------------|-----------------|-----------------|-----------------|-----------------|
| $\partial Y / \partial X_j$ | 8 th | | 3 rd | 4 th | | 7 th | | 6 th | | 5 th | 2 nd | 1 st |
| $S_{\{j\}}$ | | 7 th | 2 nd | 8 th | 5 th | | 4 th | | 6 th | 3 rd | | 1 st |
| $S_{\{j\}}^{\text{tot}}$ | | 3 rd | 1 st | 2 nd | 7 th | | 4 th | | 8 th | 5 th | | 6 th |

We can normalize local derivatives

$$S_j^{\text{loc}} = \frac{V[X_j]}{V[Y]} \left(\frac{\partial Y}{\partial X_j} \right)^2.$$



Derivative $\frac{\partial Y_{01}}{\partial X_j}$ (left), non dimensional derivative $S_j^{\text{loc}} = \frac{V[X_j]}{\text{Var}(Y_{01})} \left(\frac{\partial Y_{01}}{\partial X_j} \right)^2$ (right, upper panel), and first-order and total Sobol' indices (right, lower panel) as functions of the number of the parameter (x -axis). The derivatives are computed for $(x_1, \dots, x_d) = (E(X_1), \dots, E(X_d))$.

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What if inputs are dependent?

Conclusion, perspectives

Why is the independent framework not always the right one?

Let us come back to the example of agro-climatic model for the water status management of vineyard.

The soil texture was initially described by 3 scalar parameters: the percentages of **argil**, **sand** and **silt**.

These parameters are not independent as

$$\% \text{ argil} + \% \text{ sand} + \% \text{ silt} = 100\% .$$

In the study, this set of parameters has been replaced by a unique parameter `aSoil` describing the influence of the soil texture on its evaporation capacity.

Daily precipitations, **solar radiation**, **mean air temperature** and **potential evapotranspiration** are **temporal correlated inputs**.

We chose to use kind of scenario approach: it consists in grouping the 4 temporal inputs into a single input factor, defining **a weather scenario**.

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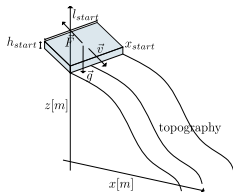
Sometimes, dependencies are due to a more complex simulation setting and cannot be handled by grouping inputs or with a scenario approach.

A [snow avalanche model](#), joint work with INRAE (Grenoble, FRANCE)

Model based on depth-averaged Saint-Venant equations (see Heredia *et al.*, 2022 for more details)

$$\frac{\partial h}{\partial t} + \frac{\partial hv}{\partial x} = 0$$

$$\frac{\partial hv}{\partial t} + \frac{\partial}{\partial x} \left(hv^2 + \frac{h^2}{2} \right) = h(g \sin \theta - F)$$



with $v = \|\vec{v}\|$ the flow velocity, h the flow depth, θ the local angle, t the time, g the gravity constant and $F = \|\vec{F}\|$ a frictional force. The model uses the Voellmy frictional force $F = \mu g \cos \theta + g / (\xi h) v^2$, where μ and ξ are friction parameters.

Equations are solved with a finite volume scheme Naaim *et al.* (98) .
The topography is the one of a path located in Bessans, France.

Let us present one of the two scenarii presented in Heredia *et al.* (2022).

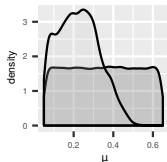
| Input | Description | Distribution |
|--------------------|---|---------------------------|
| μ | Static friction coefficient | $\mathcal{U}[0.05, 0.65]$ |
| ξ | Turbulent friction [m/s^2] | $\mathcal{U}[400, 10000]$ |
| l_{start} | Length of the release zone [m] | $\mathcal{U}[5, 300]$ |
| h_{start} | Mean snow depth in the release zone [m] | $\mathcal{U}[0.05, 3]$ |
| x_{start} | Release abscissa [m] | $\mathcal{U}[0, 1600]$ |

Let's $\text{vol}_{\text{start}} = l_{\text{start}} \times h_{\text{start}} \times 72.3 / \cos(35^\circ)$ instead of h_{start} and l_{start} .

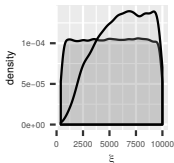
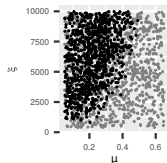
AR rules:

- ▶ avalanche simulation is flowing in $[1600\text{m}, 2412\text{m}]$,
- ▶ $\text{vol} > 7000\text{m}^3$,
- ▶ runout distance $< 2500\text{m}$ (end of the path).

From $n_0 = 100\,000$ initial runs, we keep $n_1 = 6152$ constrained ones.



correlation
original / AR
0/0.31



An alternative, the Shapley effects

We define

$$\phi_i = \frac{1}{d} \sum_{\mathbf{u} \subseteq -\{i\}} \binom{d-1}{|\mathbf{u}|}^{-1} (\text{val}(\mathbf{u} + i) - \text{val}(\mathbf{u}))$$

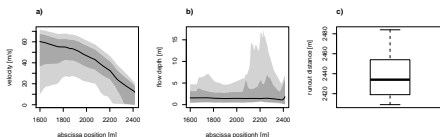
with the characteristic function $\mathbf{u} \mapsto V[E[Y|\mathbf{X}_{\mathbf{u}}]]/V[Y]$. The ϕ_i s have been introduced as the Shapley effects in [Owe14].

Interpretation

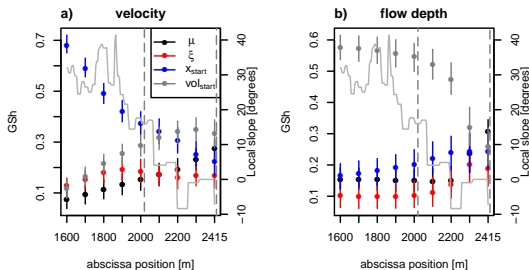
If we consider the inputs X_1, X_2, \dots, X_d as the team members trying to explain the variance of the output Y , then the set of Shapley effects $\{\phi_1, \dots, \phi_d\}$ is the unique way to allocate $V[Y]$ to all players characterized by desirable properties known as Shapley axioms (see [Sha53] for more details).

What if inputs are dependent?

How does it look like for the avalanche application?



Aggregated Shapley effects of velocity and flow depth curves calculated over space intervals $[x, 2412m]$ where $x \in \{1600m, 1700m, \dots, 2412m\}$



We have $n = 6152$, $N_{tot} = 2002$, $B = 500$. Effects are estimated using the first (2, *resp.* 4) fPCs (see Yao *et al.*, 2005, Ramsay *et al.*, 2005) explaining more than 95% of the variance. Local slope is drawn with a gray line. A gray dotted rectangle is drawn at $[2017m, 2412m]$ where avalanche return periods vary from 10 to 10 000 years.

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Conclusion, perspectives

UQ is an essential phase of forecasting.

SA helps:

- ▶ in understanding model behavior,
- ▶ as a preliminary step to model calibration,
- ▶ as a tool for decision support.

It may help in many other tasks:

- ▶ it is possible to use GSA for the construction of ensembles based on parameter perturbation (see Meryl Wimmer's PhD, 2021);
- ▶ it is possible to combine a GSA and a recursive Bayesian filtering approach for data-driven data assimilation (see, e.g., Hirvoas *et al*, 2022).

Important issue: how to perform UQ with as few as possible model runs?

Even so, SA from initial model can be **prohibitive**. Thus the importance of **metamodelling**: Gaussian Process Regression (kriging), Polynomial Chaos, Physics-Informed Neural Networks. . .

Today I only presented the basics of SA. A deeper review with practical implementation in R can be found in



with codes freely available at

<https://bookstore.siam.org/cs23/bonus>

Thanks for your attention!

Questions?

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